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Carbon Dioxide Utilization and Sequestration in Kerogen Nanopores

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Abstract

Carbon dioxide (CO₂) has been injected into oil reservoirs to maximize production for decades. On the other hand, emitted CO₂ from industrial processes is captured and stored in geological formations to mitigate greenhouse gas effects. As such, greater attention is drawn to the potential of utilizing the captured CO₂ in EOR processes. A significant portion of the injected CO₂ remains trapped due to capillary forces and through dissolution in residual liquids. In organic-rich shales, the presence of isolated kerogen nanopores add to the sequestration process due to the adsorptive nature of the surface and its preference to CO₂ over methane (CH₄), in addition to the sealing capacities of these formations. This work summarizes the latest findings of the literature with the purpose of defining further areas of investigation to fully capitalize on the potential of CO₂ sequestration and utilization in kerogen nanopores.

Keywords: CCUS, enhanced oil recovery, organic-rich shales, anthropogenic CO₂, kerogen nanopores

1. Introduction

Carbon dioxide capture, utilization and storage (CCUS) technologies involve capturing carbon dioxide (CO₂) emissions to create a synergy between the high demand for fossil fuel and mitigating greenhouse gas effects at the lowest possible cost.

CCUS captures over 90% of CO₂ emissions from power plants and industrial facilities and is predicted to reduce global gas emissions by 14% in 2050. Bearing in mind that, fossil fuel-fired power plants in the United States account for 30% of U.S. total greenhouse gas (GHG) emissions, which will only continue to increase regardless [1]. The capacity of CO₂ utilization

and storage in the U.S. is approximately 30 billion metric tons, equivalent to 35 years of CO₂ emissions captured from 140 Gigawatts (GWs) of coal-fired power [2, 3].

The captured CO₂ emissions are usually injected into geologic formations such as deep saline aquifers for storage, but most recently associated with enhanced oil recovery (CO₂-EOR) in oil and gas reservoirs. Although, CO₂-EOR has been practiced for decades now, recent advances combine the recovery process with CO₂ sequestration.

CO₂-EOR involves the injection of CO₂ into an oil/gas reservoir to recover more hydrocarbons (oil and/ or gas). Mostly, the volume of the injected CO₂ differs from that of the produced fluid with CO₂, indicating trapping or storage. Hence, incorporating the storage of anthropogenic CO₂ into CO₂-EOR in already developed oil and gas reservoirs seems economically and technically feasible. Different forms of trapping mechanisms, such as hydrodynamic and capillary trapping hold the CO₂ in place to prevent movement/leakage, ubiquitous to almost all oil and gas reservoirs [2, 3].

The United States (US) leads the world in both the number of CO₂-EOR projects and in the volume of CO₂-EOR oil production due to complimentary geology (low thermal gradient and high permeability) in the Permian Basin, located in West Texas and southeastern New Mexico [4]. Approximately 11 trillion cubic feet (560 million metric tons) total volume of CO₂ is utilized in by US CO₂-EOR as compared to 100 trillion cubic feet (5090 million metric tons) per year of total US CO₂ emissions from industrial sources [1, 4–6].

Although, CO₂ storage during CO₂-EOR in conventional oil and gas reservoirs is proven effective, the potential to sequester in unconventional organic-rich shales (gas/oil) is even more promising and economical, yet there has been minimum attention given to these vast resources. Organic-rich shales are naturally suited for CO₂ storage due to the ultra-tight impermeable nature of the formation, which would curtail CO₂ leakage. Moreover, the adsorptive surface of kerogen and kerogen nanopores in shales can store substantial amounts of CO₂ in its adsorbed state [5–7]. Thus, in depleted shale gas reservoirs, injected CO₂ replaces methane (CH₄) in the kerogen micro and nanopores and adsorb to the kerogen surface for storage [7–9]. This chapter therefore investigates the potential of CO₂ sequestration in kerogen nanopores.

2. Carbon capture

Carbon capture technology started in the 1970s in North America at industrial projects before it was applied to power generation [1]. Early application of carbon capture on a commercial basis was focused on the removal of CO₂ as part of certain industrial processes in concentrated streams [1, 8]. The Department of Energy (DOE) estimates that approximately 30 million metric tons per year of pure CO₂ are currently produced at industrial facilities located within 50 miles of existing CO₂ pipeline networks [10].

Some industrial processes with large-scale carbon capture in commercial operation include coal gasification, ethanol production, fertilizer production, natural gas processing, refinery hydrogen production, and coal-fired power generation [7, 10].

- Natural sources of CO₂ are made up of underground accumulations of naturally occurring gases with 90% CO₂. As of 2015, the natural sources are projected to account for approximately 65 Mt/a of CO₂ [8].
- Natural-gas processing are also naturally occurring underground accumulations but with significant methane content. The contribution of natural-gas processing has increased from 5 Mt/a of CO₂ in 2000 to a projected 20 Mt/a of CO₂ in 2015 [8]. Some of the known challenges of natural-gas processing include: higher oxygen (O₂) content, lower CO₂ concentration, higher flue gas and high flame temperatures [1].
- Hydrocarbon conversion involves the conversion of crude oil (or hydrocarbon feedstock) into several (high-value) products to capture CO₂ as a by-product. This process is projected to increase to approximately 5 Mt/a based on known projects under construction and in final phase [8].

However, with the recent inclusion of power generation, new systems are designed to capture and concentrate CO₂ using the following processes [7]:

- Pre-combustion carbon capture

Fuel undergoes gasification instead of combustion to produce syngas made of carbon monoxide (CO) and hydrogen (H₂). Carbon monoxide (CO) is then converted to CO₂ through a later shift reaction, while a solvent separates the CO₂ from H₂. The pre-combustion carbon capture is mostly combined with an integrated gasification combined cycle (IGCC) power plant to burn the H₂ in a combustion turbine and the resulting exhaust heat, used to power a steam turbine [1, 6].

- Post-combustion carbon capture

It involves the use of chemical solvents to separate CO₂ from the resulting flue gas from fossil fuel combustion. This method is commonly used by modified power plants for carbon capture [7].

- Oxyfuel carbon capture

This process requires the combustion of fossil fuel in pure oxygen to render the CO₂-rich exhaust gas for capture [7].

In 2016, the US Energy Information Administration (EIA) reported that electricity generated from natural gas is expected to exceed that of coal for the first time [9]. This calls for more effective measures to be put in place to curtail greenhouse gas (GHG) effects.

2.1. Carbon capture benchmarks

There are about 21 commercial-scale carbon capture projects around the world with 22 more in development [7]. Below is a list of a few of the many benchmarks in carbon capture:

- As of 2017, the Archer Daniels Midland (ADM) Company captures CO₂ from Biofuels (ethanol) production, and stores in the Mt. Simon Sandstone, a deep saline formation, Decatur, IL. An estimated amount of 1.1 million tons of CO₂ is captured per year [1, 3].

- In 2017, the NRG Petra Nova Project, TX, captures 90% of CO₂ (approximately, 1.6 million tons of CO₂ per year) from a 240 MW slipstream of flue gas of existing WA Parish plant, and transported to a nearby oil field [7].
- In 2016, Abu Dhabi CCS Project Phase 1: Emirates Steel Industries, an operating iron and steel plant, used to capture CO₂ for enhanced oil recovery by the Abu Dhabi National Oil Company (ADNOC) [3, 5].
- In 2015, Shell Quest Project, AB, CA, a bitumen upgrader complex, captures about 1 million tons of CO₂ annually from hydrogen production units and injects it into a deep saline formation for sequestration [7].
- In 2013, Conestoga Energy Partners/Petro-Santander Bonanza Bioethanol plant, KS, an ethanol plant, captures and supplies approximately 100,000 tons of CO₂ per year to a Kansas EOR field [1, 5].
- In 2010, Occidental Petroleum's Century Plant (OPCP), TX, a natural gas processing facility, compresses and transports CO₂ stream for utilization in the Permian Basin, among others [3, 5].

3. Carbon dioxide utilization (CO₂: EOR)

CO₂-EOR has been successfully implemented for nearly half a century now to recover additional oil from developed conventional oil fields in the United States and around the world. It involves the injection of CO₂, either in its supercritical or gaseous state to re-pressurize a depleted reservoir pressure to cause residual oil held in the smaller pores by capillary forces to be released [9, 10]. CO₂, unlike other fluids, reaches miscibility with crude oil at lower pressures. Furthermore, it is less expensive than other miscible fluids. As such the injected CO₂ becomes soluble with the residual oil as light hydrocarbons from the oil dissolve in the CO₂ while the CO₂ density is high when oil contains a significant volume of light hydrocarbons [4, 11].

Upon discovery, an oil reservoir is initially produced by means of the pressure gradients within the reservoir that provides the energy to move reservoir fluids to the surface. This is called the primary production stage. Eventually, the reservoir pressure declines and flow to the wellbore ceases. At this moment, a range of secondary or tertiary (EOR) methods are implemented to recover additional volumes of oil. The primary stage only recovers about 5–20% of the original oil-in-place (OOIP), with considerable amount of oil left trapped in the pore spaces of the rock [9, 12].

The next stage of production is the secondary recovery, which involves the injection of a fluid, either gas or water to sustain and maintain the depleted reservoir drive, and simultaneously recover substantial amounts of the remaining OOIP. Treated produced water (waterflooding) is commonly used at this stage since it is less expensive and readily available. In most cases, the water bypasses the oil due to difference in viscosity leaving behind significant amounts of the remaining oil-in-place. Waterflooding results in approximately 50–60% of the OOIP

trapped, hence the need for CO₂-EOR in most oil reservoirs already replenished with water-flooding. Both primary and secondary recovery methods usually extract about 35% of the OOIP [4, 9, 10].

To produce more of the remaining oil-in-place, a tertiary oil recovery phase is implemented, where fluids (CO₂, nitrogen, enriched gas, polymer solutions or surfactant solutions) are injected to interact with the oil and cause substantial changes to the oil properties [12]. Carbon dioxide (CO₂) flooding is one of the most proven EOR methods, where CO₂ is injected either in its gaseous or supercritical state. The injected CO₂ is determined to reduce the interfacial tension, minimize the viscosity of the oil to make it lighter, cause the oil volume to swell, and eventually cause the oil to flow more freely within the reservoir to the producer wellbore [11].

CO₂ is mostly delivered to the field at a high pressure (>1200 psi) and density (5 lb./gal) into injection wells within a designed pattern based on computer simulation to optimize areal sweep of the reservoir [13, 14]. Miscibility of CO₂ with the oil is important as it causes the physical forces (interfacial tension) holding the two phases apart to disappear. It occurs at a minimum pressure (MMP), where about 95% of the OOIP is recovered. Below the minimum miscibility pressure (MMP), CO₂ and oil will no longer be miscible, the oil and gas phases separate, thereby decreasing oil production rate. Significant volumes of oil are produced during CO₂-EOR. For example, the Wasson field, a Denver unit CO₂-EOR has produced more than 120 million incremental barrels of oil through 2008, with more than 2 billion barrels of OOIP and 40% of oil remaining after Waterflooding [14, 15, 19]. All types of oil reservoirs, either carbonates or sandstone could be suitable for CO₂-EOR provided the MMP can be reached [13, 14].

The operation of a CO₂-EOR project is a closed-loop system as shown in **Figure 1**, where about half of the injected CO₂ is trapped or dissolved in the reservoir and its fluids (oil and water). The produced CO₂ with oil is separated and re-injected back into the reservoir, ensuring an increase in trapped CO₂ instead of being released to the atmosphere. In addition, CO₂-EOR provides a market and revenues for the captured CO₂ from anthropogenic (industrial and power plants) sources [14, 15]. As the project matures, the volume of injected CO₂ diminishes, while recycled volumes increase. This indicates that CO₂ is being stored in the formation through a capillary trapping mechanism [10, 13, 14].

CO₂-EOR was first tested on a large-scale in the 1970s in the Permian Basin of West Texas and southeastern New Mexico. These initial projects used separated CO₂ from processed natural gas and natural sources of CO₂ instead of anthropogenic CO₂ from industrial power plants. **Figure 2** shows CO₂-EOR projects carried out around the world and in the U.S. from the 1970s to present day [9, 16]. Three developed source fields include, Sheep Mountain in south central Colorado, Bravo Dome in northeastern New Mexico, and McElmo Dome in southwestern Colorado [10]. The recent depletion of the natural source fields of CO₂ and size limitation of the pipelines for CO₂-EOR processes have paved the way for anthropogenic supplies of CO₂. In so doing, subsequent projects employ CO₂ molecules from captured emissions to supply large quantities of CO₂ for EOR processes in oil fields. Technological advancement in CO₂-EOR applications, such as 3D seismic and geomodeling reduce the rise of failures and improves the flooding efficiency [10, 13–15].

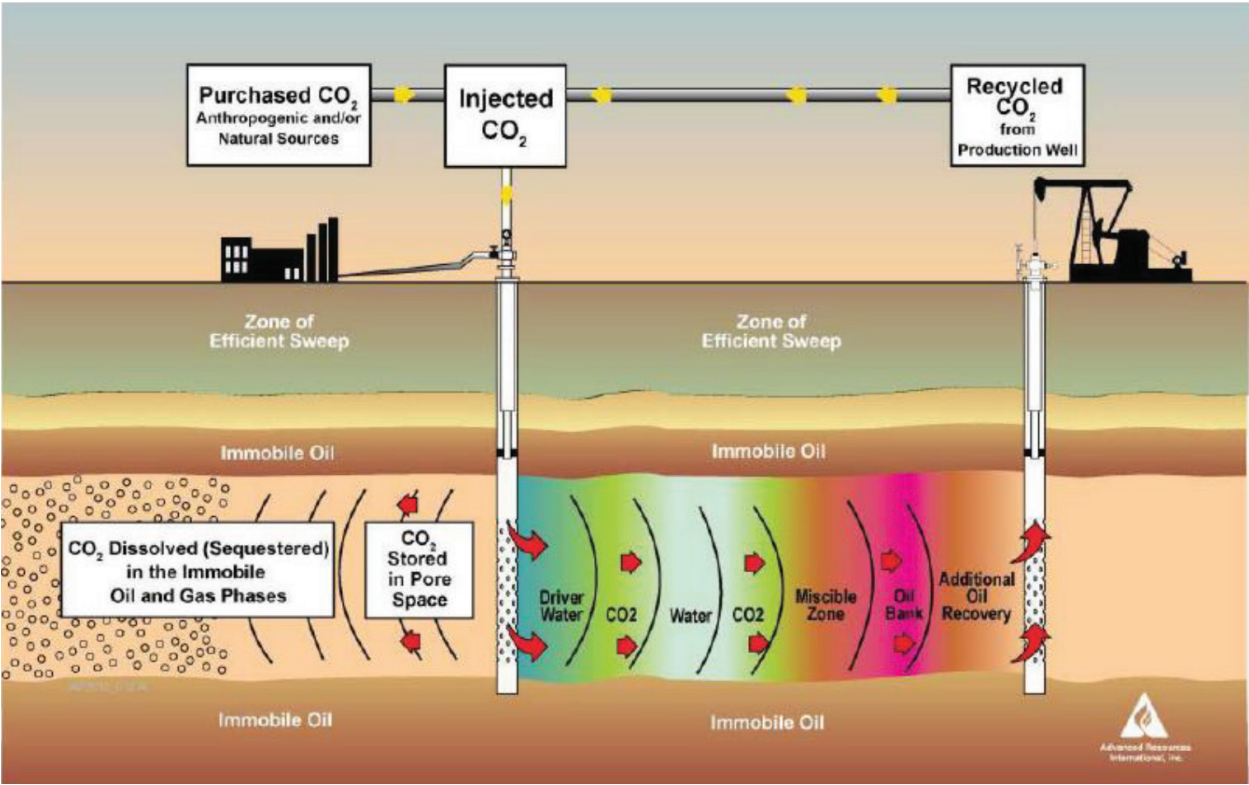


Figure 1. Schematic diagram of a closed loop CO₂-EOR [4].

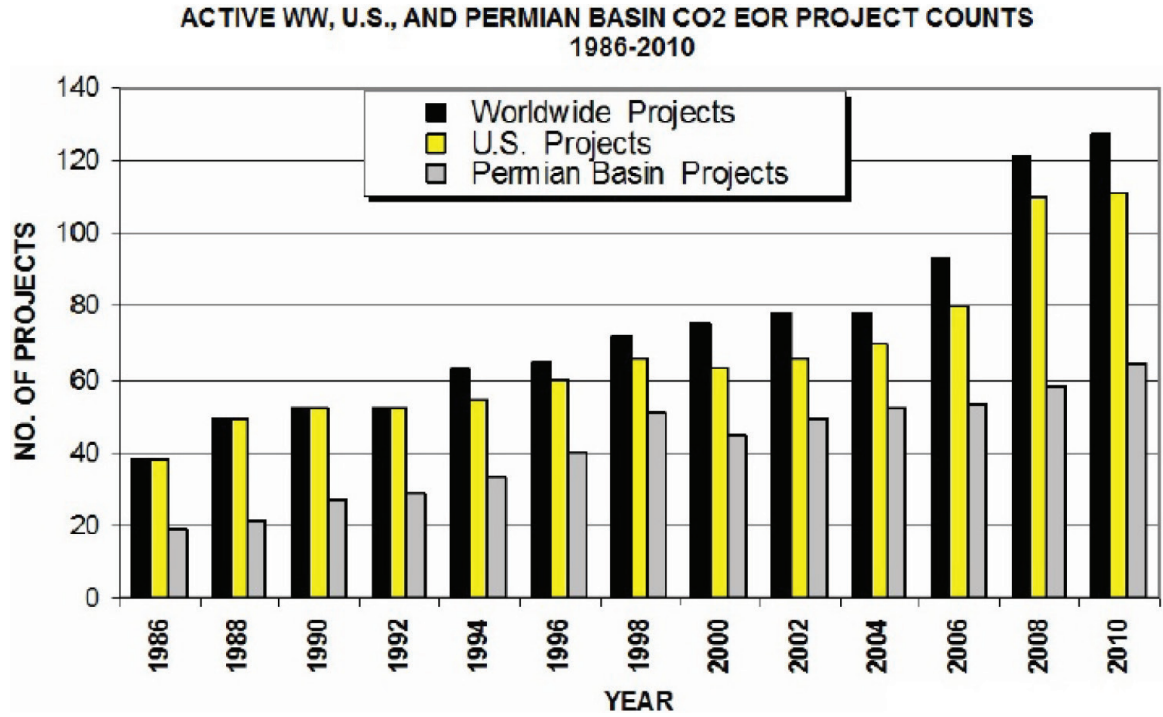


Figure 2. CO₂-EOR projects conducted worldwide and in the U.S [10].

4. Carbon dioxide sequestration

The potential of CO₂ sequestration in geologic formations is possible from the fact that certain reservoirs naturally trap and store oil and natural gas over long geological time periods until extracted [1, 4, 6, 16]. In so doing, CO₂ from power plants and industrial facilities can be trapped and stored in potential geologic formations. A large percentage of the originally injected CO₂ gets trapped in the pores of the geologic formation, while a portion of it is dissolved in the oil and also end up trapped [3, 6, 17]. These trapping processes continue as long as the CO₂ is injected. Percentage of stored CO₂ is based on total injected volumes and not on the purchased volume and is given as [10].

$$CO_{2storage}(\%) = (CO_{2injected} - CO_{2produced} - CO_{2losses}) / CO_{2purchased} \quad (1)$$

where, $CO_{2storage}$ is the CO₂ storage in metric, $CO_{2injected}$ is the total CO₂ injected, $CO_{2produced}$ is the CO₂ produced, and $CO_{2purchased}$ is the purchased CO₂ injected. CO₂ losses is estimated as the difference between total CO₂ injected and CO₂ produced. Losses may be due to leakages, infrequent power outages, among others [10].

CO₂ can be injected into conventional geological formations and stored deep underground. Most of these conventional geologic formations are at depths greater than 800 m, which eventually converts the injected CO₂ into its supercritical state. The supercritical CO₂ with a higher density than its gaseous state results in a given volume of rock capable of holding more mass of CO₂ [4, 7]. For an effective conventional geological storage, approximately 90–95% of the injected CO₂ for will be sequestered within the reservoir [4, 9, 16].

4.1. Storage mechanisms in conventional reservoirs

Trapping mechanisms encountered in CO₂-storage in conventional geologic formations include [9–11]:

- Physical trapping: hydrodynamic, stratigraphic, or structural) trapping

This involves the migration of generated hydrocarbons from organic matter (source) over long geological periods from the source rock to porous and permeable reservoir rock initially saturated with brine. The accumulated hydrocarbons are trapped below a non-permeable cap rock to prevent further migrations, and the density difference between the fluids separates the fluids into layers with gas on top, followed by oil and brine at the bottom. A similar mechanism is encountered in the case of CO₂ storage, where the less dense supercritical CO₂ plume rises due to buoyancy forces and is prevented from escaping by overlying low permeability cap rock [15]. This mechanism is considered to be relatively fast but requires characterization of the cap rock [2, 3].

- Solubility trapping

CO₂ is widely accepted to be soluble in water, as such, dissolved CO₂ can be safely stored in a geologic formation under solubility trapping. Since the CO₂-saturated brine is denser than the

unsaturated surrounding brine, density difference causes the denser brine to migrate deeper into the formation and slowly dilutes the unsaturated brine through contact. Reservoir pore pressure, temperature, and salinity of formation water are vital for solubility trapping [16]. This process occurs faster than pure diffusion, prevent CO_2 from hydrodynamically separating from other phases, and it is estimated to begin between a year and hundreds of years after CO_2 injection, also dependent on the permeability of the formation in question [2, 20].

- Mineral trapping

This process occurs over longer geological timescales than the other trapping methods, but is equally important [3]. It involves the formation of carbonic acids (H_2CO_3) as a result of CO_2 dissolution in formation brine. The resulting acid is unstable and dissociates to form groups, which react with the formation rock over long periods of time [2, 3]. In situations where, carbonate minerals are precipitated through the reaction, CO_2 is permanently trapped as a result [23].

- Capillary (residual) trapping

In a conventional sandstone oil reservoirs, brine is mostly designated as the wetting phase, while oil and gas are the non-wetting phases. In the case of carbonate rocks, oil is the wetting phase and water and gas are the non-wetting phases. In capillary trapping, the formation wetting phase surrounds the CO_2 and traps it as immobile pore scale bubbles. This process occurs over shorter time scale (right after injection) [15] compared to the other trapping mechanisms [2, 19]. In effect, the rock surface is presumed to be less water-wet in the presence of CO_2 and in the absence of oil [9].

These trapping mechanisms occur in geologic storage including [13, 14]:

- Depleted oil and gas reservoirs

Not only do these geologic formations provide a means for storing CO_2 , but also offer economic opportunities as the injected CO_2 recovers additional oil from depleted oil and gas reservoirs. Moreover, additional revenue can be obtained from the cost of selling captured CO_2 to EOR operators to fund the cost of capture technology at industrial facilities and power plants [4, 14, 17]. CO_2 is injected underground and remains immobile due to some of the enumerated trapping mechanisms listed above [3, 20].

- Deep saline formations

Saline aquifers are preferred due to their large capacities and being geographically widespread. These include porous rock formations saturated with brine at greater depths with overlying shale cap rocks, which are impermeable and act as a seal to prevent CO_2 from leaking [4, 17]. The confined CO_2 also undergoes dissolution in the brine, as well as capillary trapping to render the injected CO_2 immobile. A study [2] was carried out to measure the maximum saturation and the form of capillary curve in a CO_2 – Berea sandstone system through coreflood experiments, representative of a storage location. A capillary trapping capacity of 7.8% of the rock volume for CO_2 – Berea sandstone was recorded [2]. This is to say,

if this much is recorded in an unconsolidated formation, how much more there is to expect in a consolidated formation.

- Coal beds

Coal beds are either too deep or too thin to be economically developed, as such, they could offer CO₂ storage potential due to the adsorptive nature of the pore surfaces [4, 13]. In CO₂-enhanced coalbed methane (ECBM) production, CO₂ is injected into deep coal seams to desorb methane gas to be extracted and preferentially adsorb onto the mineral surface for permanent CO₂-storage. Yet they are not thoroughly characterized and are on a small magnitude for CO₂-storage [8, 19, 21].

4.2. Storage criterion

Nonetheless, not all geologic formations will effectively store CO₂ with minimum risks of leaking due to the buoyancy of CO₂ gas. The criteria for secure storage involve some of the following parameters (**Table 1**) as reported in a successful project carried out in Canada [10].

Table 2 summarizes CO₂-EOR and CO₂ storage projects carried out in some major oil basins around the world. A total of 1297 billion barrels of CO₂ has been utilized worldwide for CO₂-EOR, while a total of 370 billion metric tons has been stored/sequestered in the process [4].

4.3. Carbon storage regulation

CO₂ storage site selection and injection are regulated by the U.S. Federal and State agencies, in addition to checking systems for CO₂ capture and storage to reduce the potential risk of stored CO₂ to humans and the environment [1, 10, 18]. Specific regulations and particular tools are commonly implemented to selected reservoirs by different companies and agencies [1, 10].

Furthermore, the Safe Drinking Water Act (SDWA) and the U.S. Environmental Protection Agency (EPA) impose safety requirements on CO₂ injection and monitoring. Whereas, the Underground Injection Control Program (UICP) considers the previous seismic history as a requirement in selecting geologic CO₂ sequestration sites to reduce the risk of small earthquakes as well as the effect of earthquakes on leakage of CO₂. **Table 3** presents a list of monitoring tools used for CO₂-EOR and CO₂ storage projects.

Adequate depth (> 1000 meters)
Strong confining seals
Minimally faulted, fractured or folded
Adequate volume and permeability for storage
No significant diagenesis

Table 1. Criteria for storage on a basin scale [10].

Region	CO ₂ -EOR (Billion Barrels)	CO ₂ Storage capacity (Billion Metric Tons)
Asia Pacific	47	13
Central & South America	93	27
Europe	41	12
FSU	232	66
Middle East/North Africa	595	170
North America/Other	38	11
North America/U.S.	177	51
South Africa/Antarctica	74	21
TOTAL	1297	370

Table 2. CO₂-EOR and CO₂ storage in major oil basins of the world [4].

Cement integrity logs
Injection logs
Pattern and material balance techniques
Tracer injection/logging
Step rate testing
Fluid levels and reservoir pressure

Table 3. Reservoir monitoring tools used in CO₂-EOR [10].

5. CO₂ storage in unconventional shale reservoirs

As previously mentioned, conventional oil and/gas reservoirs form from the migration of petroleum and natural gas from the source (organic matter) into permeable reservoir rocks. On the other hand, unconventional shale gas/oil serve as both the source and reservoir for natural gas and liquid hydrocarbon (oil and gas condensate). These shale formations are being developed widely for oil and gas production especially in the United States (U.S) and other parts of the world. Moreover, shale formations are much more abundant and widely distributed [17] than deep un-mineable coal seams and/ or depleted oil and gas reservoirs but have not been extensively analyzed for CO₂ sequestration [19]. This is attributed to the ultra-tight nature of shales but the recent advances in horizontal drilling and hydraulic fracturing offers a new perspective into these formations [5, 19].

Shales consist of a mineral matrix (clay, pyrite, carbonate, quartz) embedded with dispersed dark kerogen (organic matter) areas as shown in **Figure 3**. Kerogen is the insoluble solid-phase nanoporous component of organic matter (decomposed plant and animal debris) in

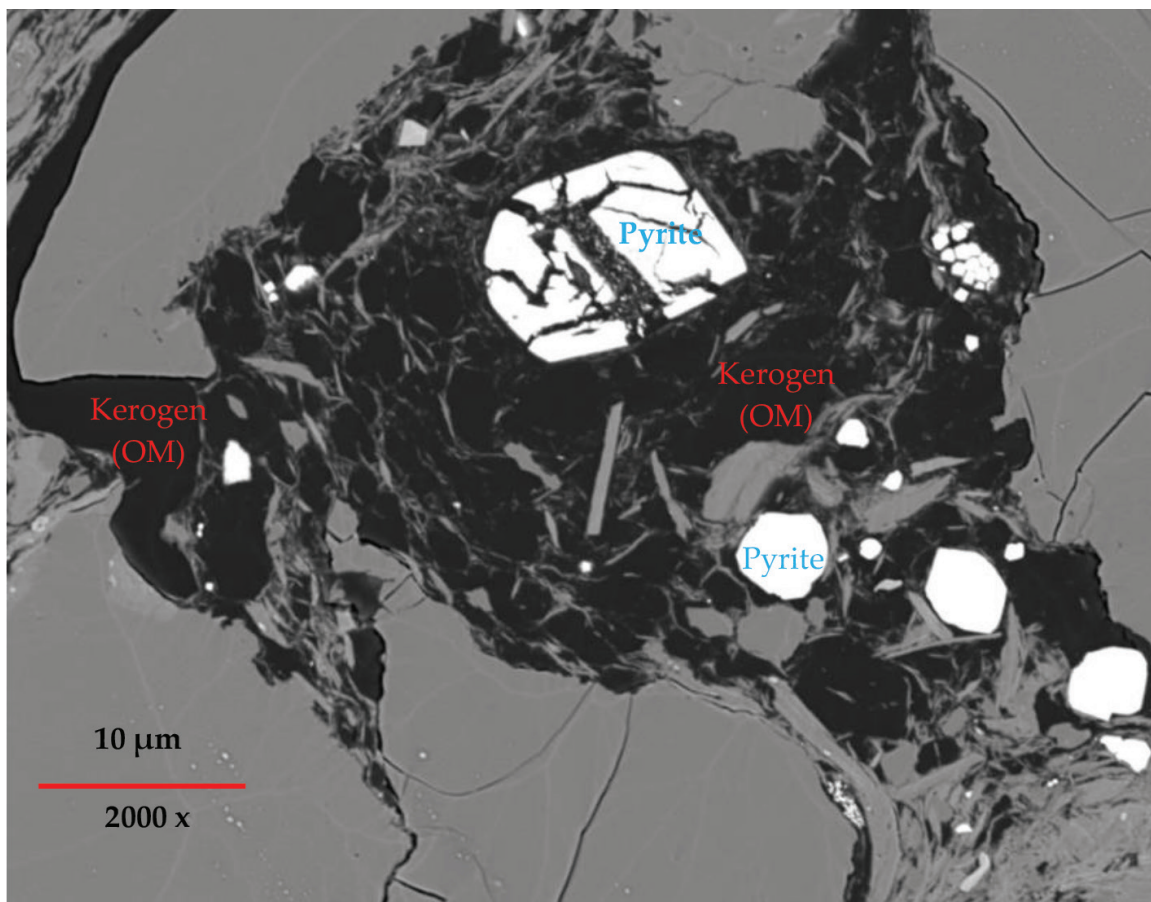


Figure 3. Backscattered electron (BSE) image of Chattanooga shale, Barber County, KS.

shale formations, which controls the gas adsorption capacity. It undergoes different stages of maturity (decomposition) at higher temperatures to produce petroleum and natural gas within the micropores (<2 nm) and mesopores (2–50 nm) [6]. The kerogen pores create a sieve for smaller CO₂ molecules, making shales more attractive for CO₂ sequestration unlike methane (CH₄) and other gas molecules [18, 30]. Thus, shales can adsorb substantial amounts of CO₂ on kerogen as well as fracture surfaces [19, 24]. The level maturity of kerogen is measured by the vitrinite reflectance (% Ro), which indicates the onset of oil (0.6–1.0 Ro%), wet gas (<0.80% Ro) and natural gas (>1.4% Ro) generations, respectively [20, 21]. Gas from shale formations are either thermogenic (generated from cracking of organic matter or the secondary cracking of oil) or biogenic (generated from microbes) [22, 26].

Because the source rock doubles as the reservoir, shales are characterized as very low permeability formations, which form strong confining seals in their own right but have surface adsorptive characteristics. As such, they require the creation of hydraulic fractures to form conduits for introducing fluids and producing them to the surface through horizontal wells. Hydraulic fracturing cracks the shale rock through injections of water, sand and chemicals at high pressure [16]. Horizontal wells with multi-stage hydraulic fractures can then be used to inject CO₂ for storage in depleted shale gas and oil reservoirs. The horizontal wells as opposed to vertical wells in conventional geologic formations add to the effectiveness of CO₂

sequestration in shales since the horizontal wells contact more of the shale formation and as a result, increase the subsurface production area of the well [19]. More so, CO₂ sequestration in shales would not require new infrastructure unlike in conventional saline aquifers [19, 20].

Most of the shale formations are located at greater depths, where the injected CO₂ is in its supercritical state, which is much preferred for both CO₂ – enhanced gas/oil recovery (EGR)/EOR in addition to CO₂ sequestration. The injected CO₂ for EGR/EOR in organic-rich shales adsorb onto the rock surface, while concurrently releasing methane gas (CH₄) and/ or oil for natural gas and oil productions, respectively [8, 22]. Furthermore, since most of the injected CO₂ would be adsorbed to the surface of kerogen rather than exist as free gas, the problem of leakage is minimized [8]. Hence, CO₂ sequestration in shales is feasible but requires knowledge of the characteristics of different shale formations as well as gas-water-rock interactions, multiphase flow, and reservoir modeling, monitoring and verification [22, 25].

Tao and Claren [19] introduced a computational method based on historical and projected methane (CH₄) production to estimate the capacity of CO₂ sequestration in Marcellus shale in eastern United States. From the results obtained, the Marcellus shale is expected to store between 10.4 and 18.4 Gt of CO₂ (approximately 50% of total US CO₂ emissions) between now and 2030. Another point to note from Tao and Claren [19] was that injected CO₂ moves through the shale formation faster than producing CH₄ through mass transfer kinetics, which enhances CO₂ sequestration process in shales. In addition, other major shale plays like Barnett, Eagle Ford, Woodford, could provide incremental storage capacity.

Nuttal [22] performed experiments to estimate CO₂ sequestration capacity in organic-rich Devonian black shales of Eastern Kentucky to be 6.8 Gt [19]. CO₂ was found to adsorb onto clay and kerogen surfaces. A direct correlation was observed between CO₂ adsorptive capacity and the total organic carbon (TOC), where CO₂ adsorption capacity increases with increasing TOC.

Kang *et al.* [6] examined shale capacity in organic-rich shales and their added advantage of allowing linear CO₂ molecules to penetrate smaller pores otherwise inaccessible to other hydrocarbon gases. Moreover, molecular interaction of CO₂ and kerogen ensures enhanced adsorption for CO₂ sequestration in shales. Injected gas (CO₂) molecules move through the shale formation through either the organics or inorganics (or both in most cases). In the organics, CO₂ dissolves into kerogen and diffuses into the kerogen nanopores, whereas, in the inorganics, CO₂ flows through irregularly shaped pores of clays, pyrite fambroids, quartz, and carbonates. Gas permeation and history-matching pressure pulse decay experiments revealed that significant amounts of CO₂ gas reached the organics through the inorganic pores.

Busch *et al.* [5] conducted diffusive transport and gas sorption experiments on shale samples. Effective diffusion coefficients increased (implies irreversible storage of CO₂) with a corresponding decrease in the concentration of bulk CO₂ volume in the sample. The decrease in bulk CO₂ volume is attributed to the dissolution of CO₂ in formation water (brine), adsorbed to clay and kerogen surfaces or undergoes geochemical reactions.

Furthermore, reservoir models can be built and used to predict viable CO₂ storage in shale reservoirs to model diffusivity, gas-water-rock interactions, and adsorption/desorption characteristics, among others. Notably, the presence of clay bound water is known to change the

gas sorption properties in coal formations so it is likely to manifest in shales as well. These phenomena could also be well understood through experimental methods [16, 19].

With these new insights, CO₂ sequestration in shale formations looks promising, however, the underlying physics of CO₂ sequestration in kerogen nanopores, where most of the sequestration takes place is much needed. A better understanding of the fluid dynamics in kerogen nanopores and predicting effective transport properties (diffusivity, permeability, etc.) is of utmost importance to practical CO₂ sequestration applications in shales. Also, it would aid in capitalizing on the full potential of CO₂ –EGR/EOR in organic-rich shales. Therefore, the application of lattice Boltzmann method (LBM) for CO₂ sequestration in kerogen nanopores focusing on the effect of adsorption was applied.

5.1. Mechanisms of CO₂ sequestration in shales

In addition to the trapping mechanisms in conventional reservoirs, organic-rich shales have an added advantage of trapping CO₂ through adsorption in the presence of kerogen [30]. Kerogen is the insoluble component of organic matter, and measured in the lab as the total organic carbon (TOC) through pyrolysis. Thus, both hydrodynamic trapping and trapping through adsorption are dependent on the wettability of CO₂ in shales.

Tao and Claren [19] developed a linear relationship between TOC and adsorption capacity using a number of published data sets as input into (Eqs. (2) and (3)), respectively for methane (CH₄) and carbon dioxide (CO₂).

$$[CH_4](cm^3/g) = 3.04 + 0.35(TOC(\%)) \quad (2)$$

$$[CO_2](cm^3/g) = 0.08 + 1.72(TOC(\%)) \quad (3)$$

The resulting plot showed the regression line of CO₂ adsorption capacity to be steeper than that of CH₄, implying that CO₂ is able to diffuse more readily than CH₄ into the porous kerogen due to its smaller molecular diameter [19]. Accordingly, we produced a TOC vs. gas adsorption capacity plot but with a focus on the level of TOC and its effect on gas adsorption capacity. Shale formations are in abundance and have diverse geologic settings throughout the U.S. (Appalachian basin, Williston basin, Illinois basin, Michigan basin, Permian basin, and Gulf Coast Region) for EOR and associated CO₂ storage [19, 23] but vary in kerogen content (TOC) and this variation in TOC has been found to impact the storage capacity of shales.

Figure 4 shows the TOC – gas adsorption capacity for a number of published TOC data [5, 19, 22, 23] ranging from low TOC (<1 wt. %), medium TOC (1 wt. % < TOC < 10 wt. %), high TOC (10 wt. % < TOC < 20 wt. %), and ultrahigh TOC (>20 wt. %) [24]. Highest CO₂ adsorption capacity is seen in the ultrahigh TOC region, followed by significant adsorption capacity in the high TOC region; the least adsorption capacity is observed at low TOC region. This implies that, the higher the kerogen content in shales, a significant amount of CO₂ sequestration is expected through adsorption trapping with subsequent production of significant amounts CH₄ displaced in the process. Therefore, conventional structural trapping becomes dominant

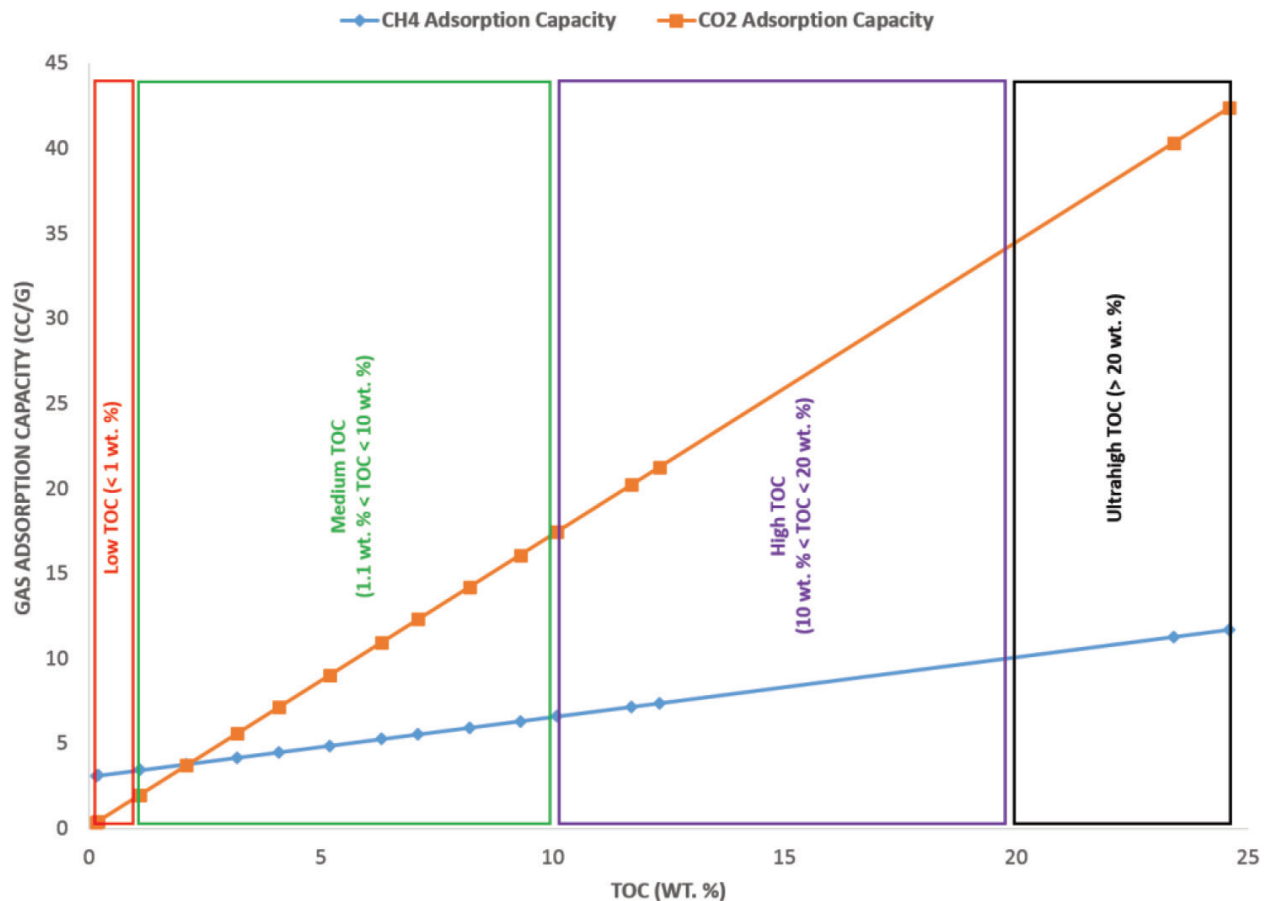


Figure 4. Gas adsorption capacity as a function of different levels of TOC.

in low TOC regions, where shale only serves as a cap rock/seal to prevent dissolved CO_2 from leaking to the surface since the mechanism of adsorption into porous kerogen (TOC) surface is close to negligible. On the other hand, within high TOC (>2 wt. %) regions, the adsorption trapping mechanism onto the kerogen surface prevails and render shale as storage medium in itself. In other words, shales with high TOC tend to be strongly CO_2 -wet, whereas shales with low TOC content exhibit water-wet conditions, with medium TOC in between strongly CO_2 -wet and water-wet conditions [24]. Furthermore, the presence of interlayering clay minerals (illite) in shales also creates a large surface area for adsorption, although the weight of TOC has a much larger influence [25].

The properties of supercritical CO_2 inside small pores are of interest for subsurface carbon storage and as such require an understanding of the processes that govern the gas transport process [19, 23]. Molecular dynamics (MD) among other microscopic computational fluid dynamics as well as analytical models based on Fick's law for gas have been applied to understand the diffusion of CO_2 and CH_4 into organic pores. While, molecular dynamics simulates kerogen pore structures with the use of molecular sieves to investigate gas transport, analytical models modify continuum approaches by incorporating slip flow and diffusion. However, molecular dynamics is not feasible to simulate gas flow in porous media at large scale due to computational time and memory constraints [25, 26] and analytical models

fall short of capturing molecular pore wall effects. On the other hand, the lattice Boltzmann method (LBM), a mesoscopic numerical method is more flexible and less time consuming since a unit of gas molecules is assigned a distribution function for simulation [25, 26].

As previously outlined, the injected gas first contacts the fracture/matrix interface and then chooses to either (1) dissolve into the organic material (kerogen) and diffuse through a nanopores network or (2) enter the inorganic material and flow through a network of irregularly shaped voids [6, 18]. Therefore, the interaction of supercritical CO₂ (scCO₂) with porous kerogen needs to be investigated for long-term reservoir storage of CO₂ in organic-rich shales. We provide a simulation study that reveals the interaction of scCO₂ with porous kerogen focusing on two key features of adsorption and diffusion.

5.2. Lattice Boltzmann simulation (LBS) of CO₂ sequestration

The lattice Boltzmann method (LBM) is a numerical method for simulating fluid at the molecular scale. This method is ideal for simulating gas flow in nanoporous kerogen since the continuum flow (Darcy's law) fails due to dominating pore-wall effects at the microscale. LBM stems from the Boltzmann kinetic theory of gases, where fluids are assumed to be made up of a large number of small particles in random motion, which undergo elastic collisions to conserve mass and momentum [19, 23, 24]. However, the LBM replaces the fluid molecules with fractious particles to reduce the number of possible particles to a handful [28]. The fractious particles are then confined to the nodes of the lattice and assigned lattice velocities (e_i) at each node as shown in **Figure 4**, where the direction index $i = 0, 1, \dots, 8$, for a D2Q9 model [29] (**Figure 5**). Following the kinetic theory, the fractious particles stream along defined lattice links and collide locally at varying lattice sites [23–25]. The streaming and collision of fluid particles by the Bhatnagar-Gross-Krook (BGK) approximation gives the lattice Boltzmann BGK equation as [19, 23–26].

$$f_i(x + e_i \Delta t, t + \Delta t) - f_i(x, t) = -\frac{1}{\tau} [f_i(x, t) - f_i^{eq}(x, t)] \quad (4)$$

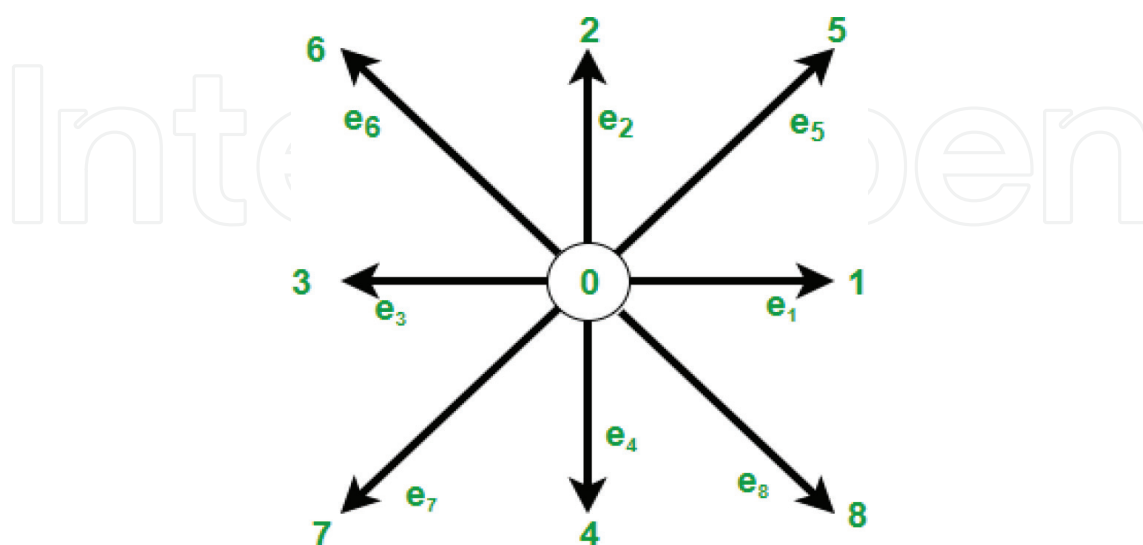


Figure 5. D2Q9 (2-D, 9-velocities) lattice nodes and velocities. Modified from [29].

where, $f_i(x, t)$ is the density distribution function, $f_i^{eq}(x, t)$ is the equilibrium distribution function, τ is the relaxation time. The left-hand side (LHS) of (Eq. (2)) represents the streaming step, while the right-hand side (RHS) constitutes the collision step.

In effect, collision of fluid particles is considered as a relaxation towards a local equilibrium, and defined for every model with varying dimensions (2-D, 3-D) and velocities (5, 9, 15, etc.).

The LBM models the distribution of and changes in the density function, from which the velocity profile is determined. Accordingly, the macroscopic fluid density and velocity are given respectively as [19, 23, 24].

$$\rho = \sum_{i=0}^8 f_i \quad (5)$$

$$u = \frac{1}{\rho} \sum_{i=0}^8 f_i e_i \quad (6)$$

where, ρ is the macroscopic fluid density and u is the macroscopic fluid velocity.

Several works [18, 29] have been carried out on modeling the convection problem encountered in deep saline aquifers during CO₂ sequestration with the lattice Boltzmann method (LBM). The findings include the fact that brine with a high CO₂ concentration was found to invade into the underlying unsaturated brine, causing an increase in the interfacial area between the CO₂-rich brine and CO₂ – deficient brine. In effect, this phenomenon enhanced the migration of CO₂ into the fracture and pores.

However, in organic-rich shales, most of the sequestration process takes place within the kerogen nanopores through adsorption [30, 31]. In so doing, there is the need to understand the interaction of supercritical CO₂ (scCO₂) with porous kerogen for long-term reservoir storage of CO₂ in organic-rich shales.

In a typical kerogen nanopore, the velocity is discontinuous at the pore wall due to the mean free path of the gas molecules exceeding the characteristic length (pore size). This phenomenon is characterized by the Knudsen number (K_n); slip flow regime falls within $0.001 < K_n < 0.1$. For chosen characteristic length 20 nm for our LBS ($K_n = 0.0243$), fluid flow falls within the slip flow regime. Slip flow boundary condition was modified for CO₂ molecules, which are predicted to not reflect at the walls but rather adsorb and desorb after some time lag [26, 27]. In effect, the velocity of the pore wall is defined to be dependent on the surface diffusion coefficient of CO₂ gas as well as Langmuir adsorption parameters based on the amount of adsorbed gas. Hence, the slip velocity at the pore-wall is given by [26, 27, 30].

$$u_{slip} = (1 - \alpha) u_g + \alpha u_w \quad (7)$$

where, u_{slip} is the slip velocity, u_g is the fluid velocity away from the wall, u_w is the local wall velocity dependent on the surface diffusion coefficient, and α is the amount of adsorbed gas at the solid surface through Langmuir isotherm.

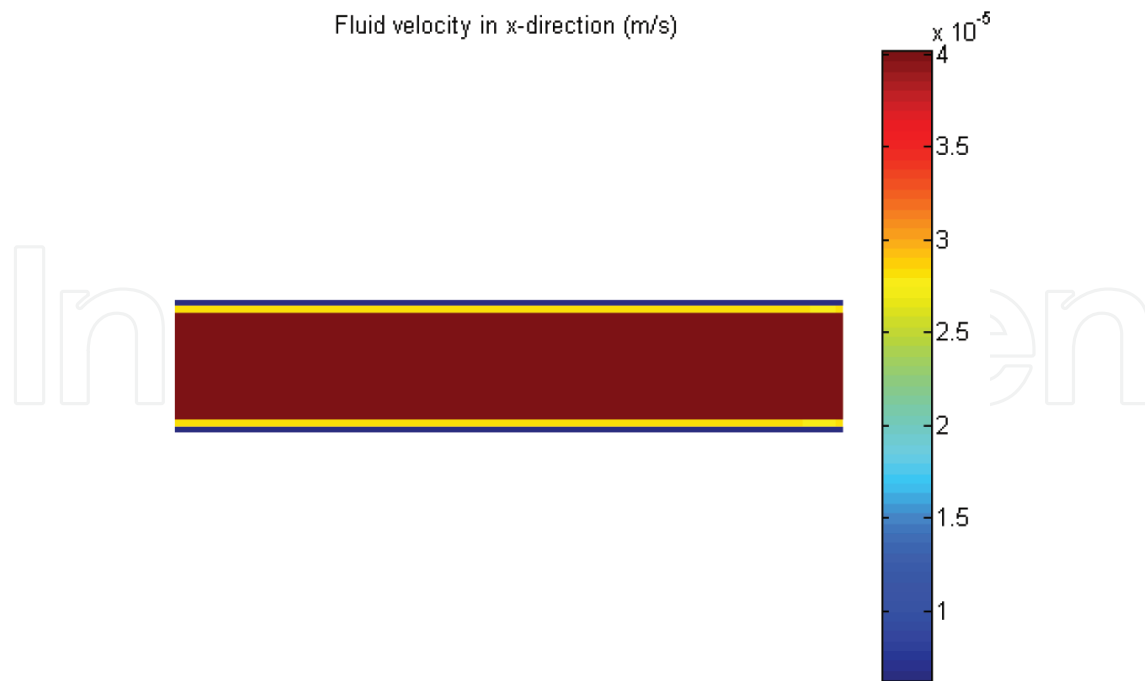


Figure 6. Composite velocity distribution of both CH₄ (center) and CO₂ (at the walls) in a 20 nm pore-slit.

Marcellus shale reservoir conditions were implemented at a high pressure of 12 MPa and temperature of 300 K. The D2Q9 LBM diffusion coefficient is known to be given in (Eq.(8)) and is directly comparable to the kinematic viscosity [28, 32].

$$D = \frac{1}{3} \left(\tau_{\sigma} - \frac{1}{2} \right) \quad (8)$$

where, D is the diffusion coefficient of the D2Q9 LBM and τ_{σ} is the relaxation time for each fluid component.

A 20-nm pore-slit is filled with both CH₄ and CO₂ at 12 MPa. Hydrodynamic velocity boundary condition is implemented at the upper and lower walls, while the pressure boundary condition is applied to the east and west ends. **Figure 6** shows the static velocity profile of both fluids in the pore-slit; CO₂ occupies the surface of the pore walls on both ends as the wetting phase, while methane occupies the center as the non-wetting phase.

Estimating the amount of adsorbed gas for CO₂ and CH₄, respectively, it was found that CO₂ adsorption capacity was much more than that of CH₄ for the same pore dimension and prevailing temperature and pressure. Furthermore, the diffusion coefficient CO₂ at $\tau = 1$ is higher than the diffusion coefficient of CH₄ at $\tau = 0.8$. The estimated magnitude of CH₄ diffusion coefficient is given in the range of 10^{-13} – 10^{-10} m²/s in carbon molecule sieves [19].

6. Conclusions

CO₂ sequestration in organic-rich shales to mitigate greenhouse gas (GHG) effects is proven to be very feasible through experimental and numerical simulations. Our literature review

and LBS suggest that organic-rich shales are capable of storing CO₂ in substantial quantities in its adsorbed state in the presence of higher TOC levels. In addition to shales being widely distributed and in abundance, the natural confining seals of the formation reduces the risk of leakage. On the other hand, CO₂-EGR/EOR can be achieved as part of the CO₂ sequestration process; CO₂-EGR/EOR produces relatively clean fuel and sustains energy demands.

However, to accurately benefit from CO₂ – sequestration in organic-rich shales, there is the need to overcome developmental challenges and understand the rock-fluid and fluid–fluid interactions in organic-rich shales for large scale pilot test and implementation.

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Conflict of interest

The authors declare no conflict of interest.

Notes

The authors declare no competing financial interest.

Appendices and nomenclature

EGR/EOR	enhanced gas recovery/enhanced oil recovery
OOIP	original oil-in-place
ECBM	enhanced coalbed methane
CO ₂	carbon dioxide
scCO ₂	supercritical CO ₂
CH ₄	methane
Gt	giga tons
TOC	total organic carbon/content

Ro %	vitrinite reflectance
LBM	lattice Boltzmann method
LBS	Lattice Boltzmann simulation
BGK	Bhatnagar-Gross-Krook
$f_i(x, t)$	velocity distribution function
$f_i^{eq}(x, t)$	equilibrium distribution function
τ	relaxation time
e_i	lattice velocities
ρ	macroscopic density
u	macroscopic velocity
GHG	greenhouse gas
MPa	mega Pascal
K	Kelvin
Kn	Knudsen number
nm	nanometer
u_{slip}	slip velocity
D	diffusion coefficient of D2Q9 LBM
D2Q9	2-dimensional, 9 velocity/speed model

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