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Ordovician and Carboniferous Volcanism/Plutonism in Central Inner Mongolia, China and Paleozoic Evolution of the Central Asian Orogenic Belt

Yuruo Shi

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Abstract

Adakite was originally proposed as a genetic term to define intermediate to high silica, high Sr/Y and La/Yb volcanic and plutonic rocks derived from melting of young, subducted lithosphere. However, most volcanic rocks in modern island arcs and continental arcs are probably derived from melting in the mantle wedge. Trace element chemistry with high Sr/Y ratios is a distinguishing characteristic of adakites. Ordovician and Carboniferous volcanic/plutonic rocks with high Sr/Y ratios occur in Central Inner Mongolia, which is situated on the southern margin of the Central Asian Orogenic Belt (CAOB). The samples are mostly granodiorite, tonalite and quartz-diorite in composition with intermediate to high-silica, high Na₂O (3.08–4.26 wt.%), low K₂O (0.89–2.86 wt.%) and high Na₂O/K₂O and Sr/Y ratios. Their chondrite-normalized REE patterns are characterized by LREE enrichment. In mantle-normalized multi-element variation diagrams, they show typical negative Nb anomalies, and all samples display positive $\epsilon_{\text{Hf}}(t)$ and $\epsilon_{\text{Nd}}(t)$ values, and low I_{Sr} . The Ordovician rocks, however, show higher Sr/Y and La/Yb ratios than the Carboniferous samples, implying that the older granitoids represent adakitic granitoids, and the Carboniferous granitoids are typical subduction-related arc granitoids but also with adakite-like compositions. The results are compatible with the view that the Central Asian Orogenic Belt (CAOB) in Inner Mongolia evolved through operation of several subduction systems with different polarities: an early-middle Paleozoic subduction and accretion system along the northern margin of the North China Craton and the southern margin of the Mongolian terrane, and late Paleozoic northward subduction along the northern orogen and exhumation of a high-pressure metamorphic terrane on the northern margin of the North China Craton.

Keywords: Adakitic, Ordovician and Carboniferous, Geochemistry, Hf-in-zircon isotopes, Central Inner Mongolia, CAOB

1. Introduction

It is generally agreed that the Solonker suture zone represents the southernmost termination of the Central Asian Orogenic belt (CAOB; [1–5]). However, there are a lot of controversies about the timing of the amalgamation of the Central Asian Orogenic belt with continental blocks to the south [1–9]. It is still debated whether the CAO evolved through subduction and accretion of a single, long-lasting, subduction system [10] or through several subduction systems with different polarities and through collision/accretion of arcs and microcontinents [11–15].

Adakite was originally proposed as a genetic term to define intermediate to high-silica, high Sr/Y and La/Yb volcanic and plutonic rocks derived from melting of young, subducted lithosphere [16]. However, most volcanic rocks in modern island arcs and continental arcs are probably derived from melting in the mantle wedge [17]. Trace element chemistry with high Sr/Y ratios is a distinguishing characteristic of adakites [16, 18]. Ordovician and Carboniferous volcanic/plutonic rocks with high Sr/Y ratios occur in Central Inner Mongolia, which is situated on the southern margin of the Central Asian Orogenic Belt (CAOB, [19]). Early Paleozoic [6–9, 20–22] and Late Paleozoic [2–4, 23] arc volcanism/plutonism as part of trench-island arc-basin systems occurred along the southern margin of South Mongolian microcontinent and the northern margin of North China Craton, suggesting concurrent two-way subduction towards opposing continental margins. The chapter focuses on early and late Paleozoic volcanic/plutonic rocks with high Sr/Y ratios in Central Inner Mongolia, and contributes geochemical data to the evolution of the CAO.

2. Geotectonic situation

Central Asian Orogenic Belt (CAOB, [19]) is a giant accretionary orogen [15], bounded by the Siberian, Tarim and North China Craton ([19, 24]; **Figure 1**), and reflects a complex evolution from the late Mesoproterozoic to late Palaeozoic [1, 6, 8, 14, 26, 27].

In Central Inner Mongolia and adjacent southern Mongolia, the Solonker suture zone can be traced for ca. 1000 km by dismembered ophiolite fragments (**Figure 1**) and represents a major paleo-plate boundary in Central Asia that stretches northeastwards for more than 2500 km in Mongolia and China [28]. It has been variably interpreted as the southernmost limit of the Altai ([10]) or the southernmost termination of the CAO [1]. The Solonker suture zone separates two continental blocks (**Figure 1**) [3]. The Northern Block consists of the Southern Mongolia (or Hutag Uul) block (gneissic granite, 1784 ± 7 Ma, Shi et al., unpublished data) and the Northern Orogen, which includes metamorphic complex (an orthogneiss has a zircon age of 437 ± 3 Ma, [29]), an ophiolitic mélangé with blueschist, a near-trench granitoid (ca. 498–461 Ma) and a juvenile arc (ca. 484–469 Ma, [3]). The Southern Block comprises the southern orogen and the northern margin of the North China Craton.

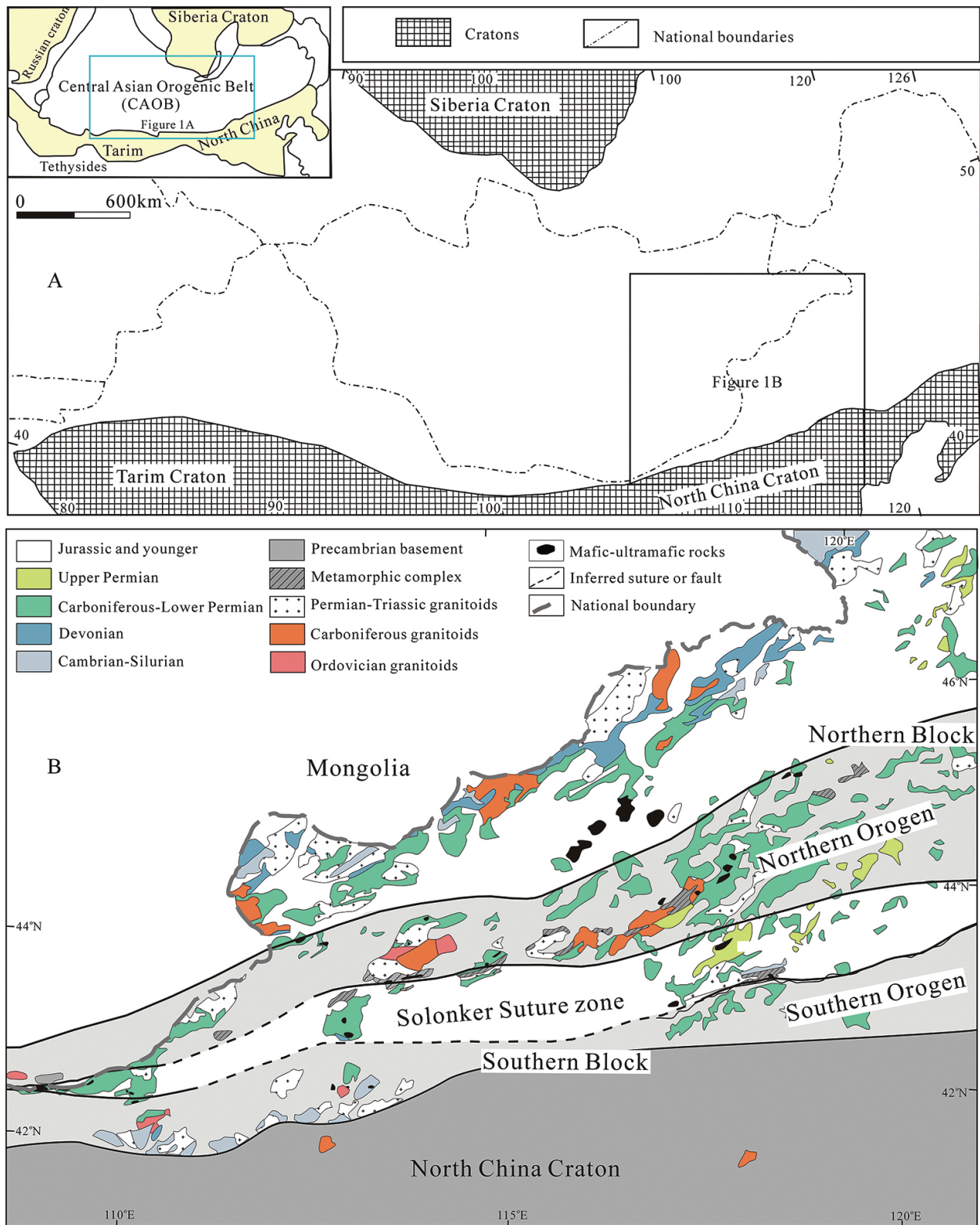


Figure 1. Geological sketch map of the southeastern CAOB (the inset map of **Figure 1A** compiled after [19]; **Figure 1B** after [3, 25]). In **Figure 1B**, the Solonker suture zone represents the tectonic boundary between the northern and the southern continental blocks [3].

Paleozoic volcanic rocks and granitoids are widely distributed along the margin of the Solonker suture zone. Ordovician granitoids (quartz-diorite, granodiorite, diorite, tonalite, and trondjemite; **Table 1** and **Figure 2**) occur in the northern and southern orogen [7, 8, 20, 21, 42, 43]; **Figure 1**), whereas Carboniferous volcanic rocks and granitoids (quartz-diorite, granodiorite, tonalite, and granite; **Table 1** and **Figure 2**) are mainly distributed in the northern orogen ([2, 23, 30, 31, 35, 37, 38, 40, 41]; **Figure 1**), and scattered along the northern margin of the North China Craton [44, 45]. The geochemical data of representative rocks are listed in **Table 2**.

Unit	Episode	Lithology	Zircon age (Ma)	Method	$\epsilon_{Hf(t)}$ (Zircon)	$\epsilon_{Nd(t)}$ (Whole rock)	Initial $^{87}Sr/^{86}Sr$ (whole rock)	Reference
Northern Orogen	Ordovician	Quartz diorite	490 ± 8	SHRIMP				[23]
		Tonalite	479 ± 8	SHRIMP		+1.5	0.7053	[7]
		Quartz diorite	475 ± 6	SHRIMP				[7]
		Granodiorite	472 ± 3	SHRIMP	+7.4 to +10.7	+2.2	0.7060	This study
	Carboniferous	Tonalite	464 ± 8	SHRIMP		+1.4	0.7053	[7]
		Tonalite	329 ± 3	SHRIMP		+5.1	0.7043	This study
		Quartz diorite	325 ± 3	SHRIMP				[30]
		Quartz diorite	323 ± 4	SHRIMP				[31]
		Quartz diorite	322 ± 3	SHRIMP				[30]
		Monzogranite	322 ± 1	LA-ICP-MS	+10.6 to +14.0			[32]
		Quartz diorite	320 ± 3	SHRIMP	+8.1 to +12.3	+2.1	0.7051	This study
		Granodiorite	320 ± 8	SHRIMP		+1.0	0.7055	This study
		Andesite	320 ± 7	SHRIMP				[33]
		Granite	319 ± 4	LA-ICP-MS				[34]
		Granodiorite	319 ± 3	SHRIMP				[35]
		Basalt	318 ± 3	LA-ICP-MS				[4]
		Granite	317 ± 2	LA-ICP-MS				[36]
		Garnet bearing granite	316 ± 3	SHRIMP				[29]
		Granodiorite	316 ± 1	LA-ICP-MS	+3.0 to +12.6			[32]
		Quartz diorite	315 ± 4	SHRIMP				[31]
		Basalt	315 ± 4	LA-ICP-MS				[4]
		Monzonitic granite	314 ± 2	LA-ICP-MS				[37]
		Quartz diorite	313 ± 5	SHRIMP				[31]
		Granodiorite	312 ± 1	LA-ICP-MS				[38]

Unit	Episode	Lithology	Zircon age (Ma)	Method	$\epsilon_{\text{Hf}(t)}$ (Zircon)	$\epsilon_{\text{Nd}(t)}$ (Whole rock)	Initial $^{87}\text{Sr}/^{86}\text{Sr}$ (whole rock)	Reference
Mongolia Hutag Uul		Monzonitic diorite	312 ± 4	SHRIMP				[39]
		Diorite	311 ± 2	SHRIMP				[35]
		Gabbroic diorite	310 ± 5	SHRIMP	+5.4 to +11.5	+2.5	0.7052	[2]
		Quartz diorite	310 ± 2	SHRIMP				[35]
		Volcanic rock	310 ± 1	LA-ICP-MS				[38]
		Quartz diorite	309 ± 8	SHRIMP		-0.2	0.7056	[23]
		Volcanic rock	309 ± 2	LA-ICP-MS				[40]
		Monzonitic granite	308 ± 2	LA-ICP-MS				[36]
		Monzonitic granite	307 ± 2	SHRIMP				[41]
		Volcanic rock	307 ± 6	LA-ICP-MS				[40]
		Rhyolite	303 ± 6	SHRIMP				[33]
		Gneissic granite	1784 ± 7	SHRIMP				Shi et al., unpub lished
		Granodiorite	454 ± 10	SHRIMP				Shi et al., unpub lished
Southern Orogen	Ordo- vician	Tonalite	491 ± 8	SHRIMP		+5.2	0.7047	[8]
		Diorite	472					[42]
		Dacite	459 ± 8	SHRIMP				[43]
		Dacite	458 ± 3	SHRIMP		+7.1	0.7058	[8]
		Quartz diorite	454 ± 4	SHRIMP		+2.0	0.7056	[8]
		Diorite	452 ± 3	SHRIMP				[8]
		Trondjemite	451 ± 7	SHRIMP				[43]
		Granodiorite	450					[42]
Northern margin of NCC	Carboni- ferous	Biotite K- feldspar granite	342 ± 5	SHRIMP				[44]
		Quartz diorite	324 ± 6	SHRIMP				[45]
		Quartz diorite	311 ± 2	SHRIMP				[45]
		Granodiorite	310 ± 5	SHRIMP				[45]
		Quartz diorite	302 ± 4	SHRIMP				[45]
Ophiolitic block								
Erlianhot- Hegenshan		Gabbro	354 ± 7	SHRIMP		+9.8	0.7043	[3]
		Gabbro	298 ± 9	SHRIMP		+8.1	0.7037	[25]
Jiaoqier-		Gabbro	483 ± 2	SHRIMP				[8]

Unit	Episode	Lithology	Zircon age (Ma)	Method	$\epsilon_{Hf(t)}$ (Zircon)	$\epsilon_{Nd(t)}$ (Whole rock)	Initial $^{87}Sr/^{86}Sr$ (whole rock)	Reference
Xilinhote								
Solonker-Linxi		Trondjemite	324 ± 3	SHRIMP		+8.4	0.7039	[3]
		Plagiogranite	288 ± 6	SHRIMP		+7.8	0.7039	[3]
		Gabbro	284 ± 4	SHRIMP		+6.8	0.7043	[3]
Wenduermiao-Xar Moron		Gabbro	480 ± 3	SHRIMP		+9.2	0.7059	[8]

Table 1. Summary of zircon ages, Hf isotopic data and whole-rock Sr-Nd isotopic data.

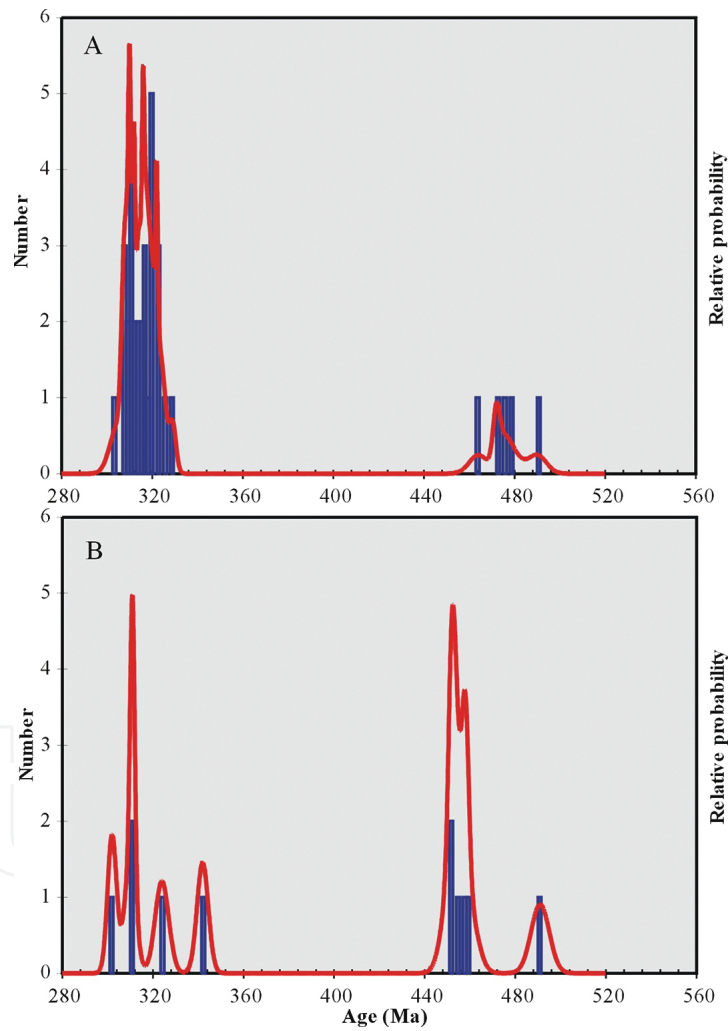


Figure 2. Cumulative plot for zircon U-Pb ages of Ordovician and Carboniferous rocks from Central Inner Mongolia (data and references are in Table 1). A for rocks from the Northern Block, which consists of the Southern Mongolia (or Hutag Uul) block and the northern orogen; and B for rocks from the Southern Block, which is composed of the northern margin of North China Craton and the southern orogen [3].

Sample	MS02-7	MB1-3	MS3-5	MB1-6	MB1-1	MB1-5	MB1-2	MB1-4
Lithology	Tonalite	Granodiorite	Tonalite	Tonalite	Granodiorite	Quartz-diorite	Granite	Granite
Age (Ma)	479 ± 8	472 ± 3	464 ± 8	329 ± 3	ca. 320	320 ± 3	297 ± 2	--
SiO ₂	61.13	67.37	61.62	61.98	66.47	54.96	75.19	71.94
TiO ₂	0.42	0.25	0.41	0.59	0.43	0.68	0.18	0.17
Al ₂ O ₃	17.05	16.31	16.56	16.22	15.63	18.80	13.76	15.68
TFe ₂ O ₃	5.88	3.68	5.62	5.82	4.17	7.98	1.92	1.30
MnO	0.14	0.08	0.14	0.08	0.06	0.12	0.02	0.02
MgO	2.34	1.14	2.27	2.95	1.69	3.65	0.70	0.47
CaO	5.69	3.91	5.78	4.93	3.43	6.25	0.39	1.54
Na ₂ O	3.56	4.26	3.08	3.22	3.37	3.14	3.92	5.47
K ₂ O	1.34	1.37	1.74	1.49	2.86	0.89	2.62	2.89
P ₂ O ₅	0.19	0.12	0.18	0.16	0.14	0.22	0.09	0.11
LOI	1.84	1.40	2.78	2.48	1.81	2.95	1.32	0.73
TOTAL	99.58	99.89	100.18	99.92	100.06	99.64	100.11	100.32
Na ₂ O/K ₂ O	2.66	3.11	1.77	2.16	1.18	3.53	1.50	1.89
Sc	12.8	5.60	13.4	17.1	9.5	20.0	2.69	0.60
V	115	64	107	123	83	151	33.9	25.1
Cr	20.44	5.0	79.5	45	21	26	8.8	8.1
Co	12.1	6.0	10.6	17.3	10.7	21.8	3.23	2.80
Ni	10.3	3.6	18	29.1	11.4	19.8	7.0	3.9
Cu	5.4	9.1	6.2	38.7	10.3	51.4	23.0	15.8
Zn	51.5	39.2	49.5	55.9	42.7	87.8	18.3	30.9
Ga	16.5	17.4	16.3	16.9	16.1	19.4	12.7	18.1
Ge	1.38	1.48	1.42	1.45	1.17	1.28	1.21	0.78
Rb	32.26	51.1	42.09	69.08	96.9	24.05	99.6	66.5
Sr	649	711	604	304	373	473	198	581
Zr	84.8	78.3	81.9	149.4	171	52.6	75.1	104.4

Sample	MS02-7	MB1-3	MS3-5	MB1-6	MB1-1	MB1-5	MB1-2	MB1-4
Lithology	Tonalite	Granodiorite	Tonalite	Tonalite	Granodiorite	Quartz-diorite	Granite	Granite
Age (Ma)	479 ± 8	472 ± 3	464 ± 8	329 ± 3	ca. 320	320 ± 3	297 ± 2	--
Nb	3.64	4.73	3.2	5.23	6.20	4.96	6.12	1.96
Cs	0.55	0.794	0.8	1.25	1.12	1.54	1.82	1.27
Ba	685.0	471.8	862.4	241.8	687.1	173.8	487.8	511.8
Hf	2.43	2.09	2.43	3.59	4.50	1.33	2.18	2.88
Ta	0.24	0.26	0.23	0.45	0.55	0.25	0.61	0.19
Th	3.78	10.46	2.76	5.83	11.26	0.49	11.83	2.31
U	1	1.48	1.05	1.277	2.16	0.264	0.58	1.11
La	10.92	25.5	7.35	14.67	19.3	9.45	8.91	5.41
Ce	22.7	49.3	17.01	29.3	37.9	20.9	28.4	14.3
Pr	2.85	4.58	2.11	3.78	4.23	2.63	1.97	1.25
Nd	11.85	15.8	9.03	16.2	16.2	11.6	7.12	5.11
Sm	2.67	2.19	2.32	3.63	3.09	2.61	1.33	1.00
Eu	0.84	0.61	0.77	1.00	0.81	0.88	0.32	0.38
Gd	2.57	2.09	2.26	3.84	2.81	2.72	1.45	0.93
Tb	0.42	0.23	0.38	0.64	0.40	0.42	0.22	0.11
Dy	2.39	1.32	2.26	3.77	2.31	2.54	1.46	0.72
Ho	0.52	0.26	0.51	0.87	0.48	0.53	0.33	0.12
Er	1.52	0.78	1.39	2.32	1.29	1.41	0.95	0.35
Tm	0.25	0.12	0.23	0.39	0.20	0.22	0.16	0.038
Yb	1.64	0.95	1.61	2.61	1.42	1.50	1.18	0.37
Lu	0.27	0.15	0.28	0.44	0.24	0.24	0.18	0.028
Y	16.2	9.1	14.5	27.0	14.1	16.9	11.1	4.68
La/Yb	6.7	26.8	4.6	5.6	13.6	6.3	7.6	14.6
Sr/Y	40	78	42	11	27	28	18	124

Table 2. Major oxide (wt.%) and trace element (ppm) composition of representative samples.

Figure 3 shows the photographs of field occurrences and photomicrographs of some representative samples. **Figure 3A** was taken from Central Inner Mongolia to show the beautiful landscape; **Figure 3B** shows the Carboniferous volcanic rocks which are located in the Southern Block.

Granodiorite sample MB1-3 (**Figure 3C** and **3D**), collected from Baiyinbaolidao, southern Sonidzuoqi, which is located in the Northern Block, is medium-grained, foliated and consists of plagioclase (45–50 vol.%), quartz (20–25%), K-feldspar (10–15), biotite (5–10%), hornblende (1–5%), accessory zircon, apatite and sphene. Plagioclase is partially epidotized, sericitized and biotite grains are chloritized.

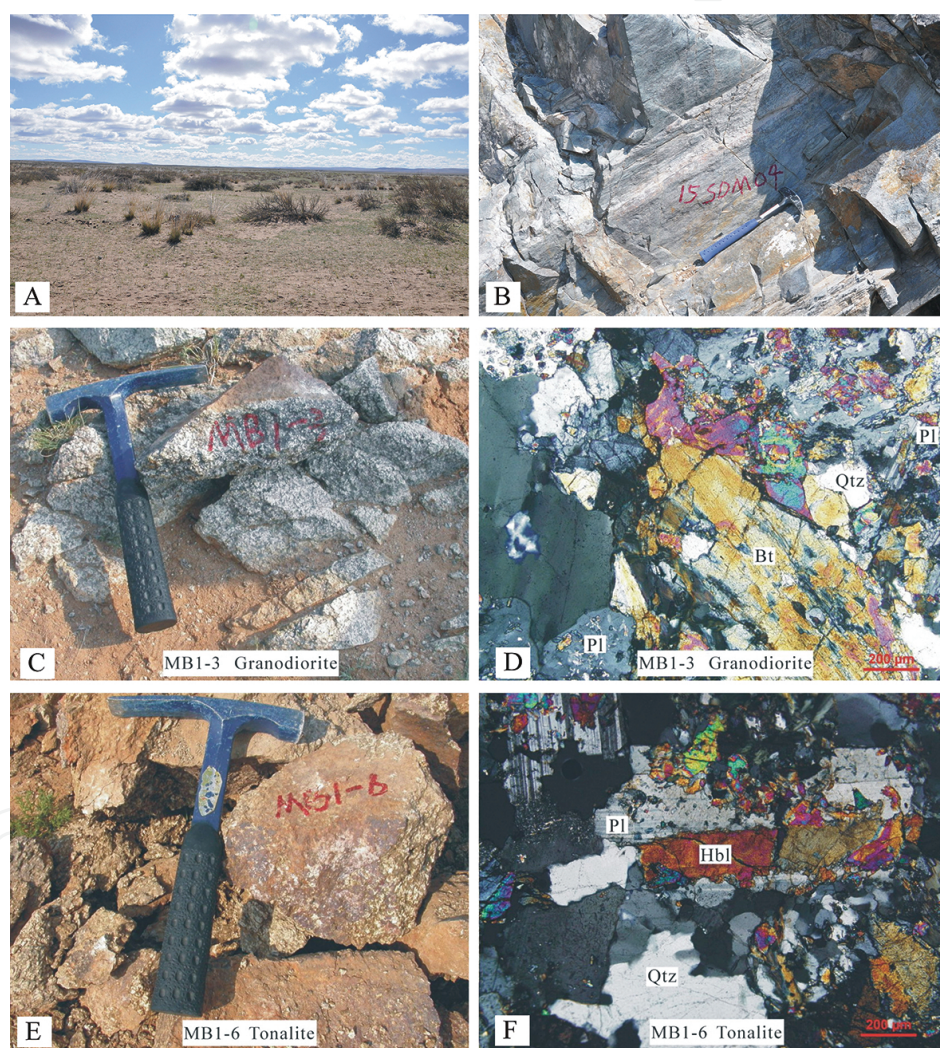


Figure 3. Photographs to show field occurrences and photomicrographs of some representative samples.

Tonalite sample MB1-6 (**Figure 3E** and **3F**), which is also collected from Baiyinbaolidao, Southern Sonidzuoqi, is medium-grained and consists of plagioclase (60–65%), quartz (20–25%), hornblende (10–15%) and biotite (1–5%) with trace amounts of zircon, apatite and sphene. Plagioclase is partially epidotized, and biotite grains are chloritized.

3. Petrogenesis of the Ordovician and Carboniferous volcanic rocks and granitoids

The Ordovician granitoid samples have intermediate to high-silica (61.13–67.37 wt.%), high Al_2O_3 (mostly >15 %), higher Na_2O than K_2O ($\text{Na}_2\text{O} > \text{K}_2\text{O}$, $\text{Na}_2\text{O}/\text{K}_2\text{O} = 1.77\text{--}3.11$), low MgO (<3%), low HREE (**Figure 4**), depleted HFSE (**Figure 5**), Y and Yb ($\text{Y} < 18$ ppm, $\text{Yb} < 1.9$ ppm), high Sr (604–711 ppm), Sr/Y mostly >40 (40.1–78.1) (**Table 2**; **Figure 6**) and low I_{Sr} with positive $\varepsilon_{\text{Nd}}(t)$ isotope ratios (**Table 1**; **Figure 7**). The Ordovician granitoid samples therefore represent adakitic compositions ([16, 48]; **Table 3**).

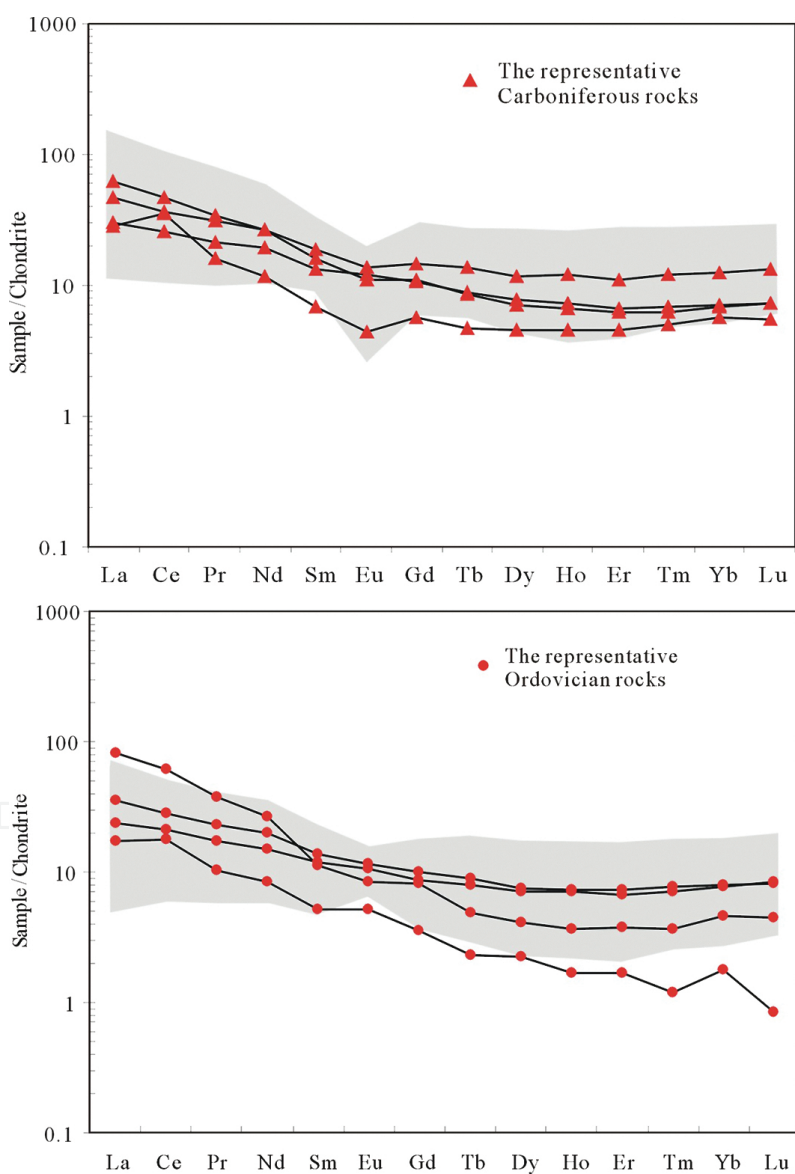


Figure 4. Chondrite (CHON)-normalized REE patterns for representative samples (grey fields show data from [7, 8, 43] for Ordovician granitoids; and from [23, 30, 32, 34, 36, 37] for Carboniferous granitoids). Chondrite values are from [46].

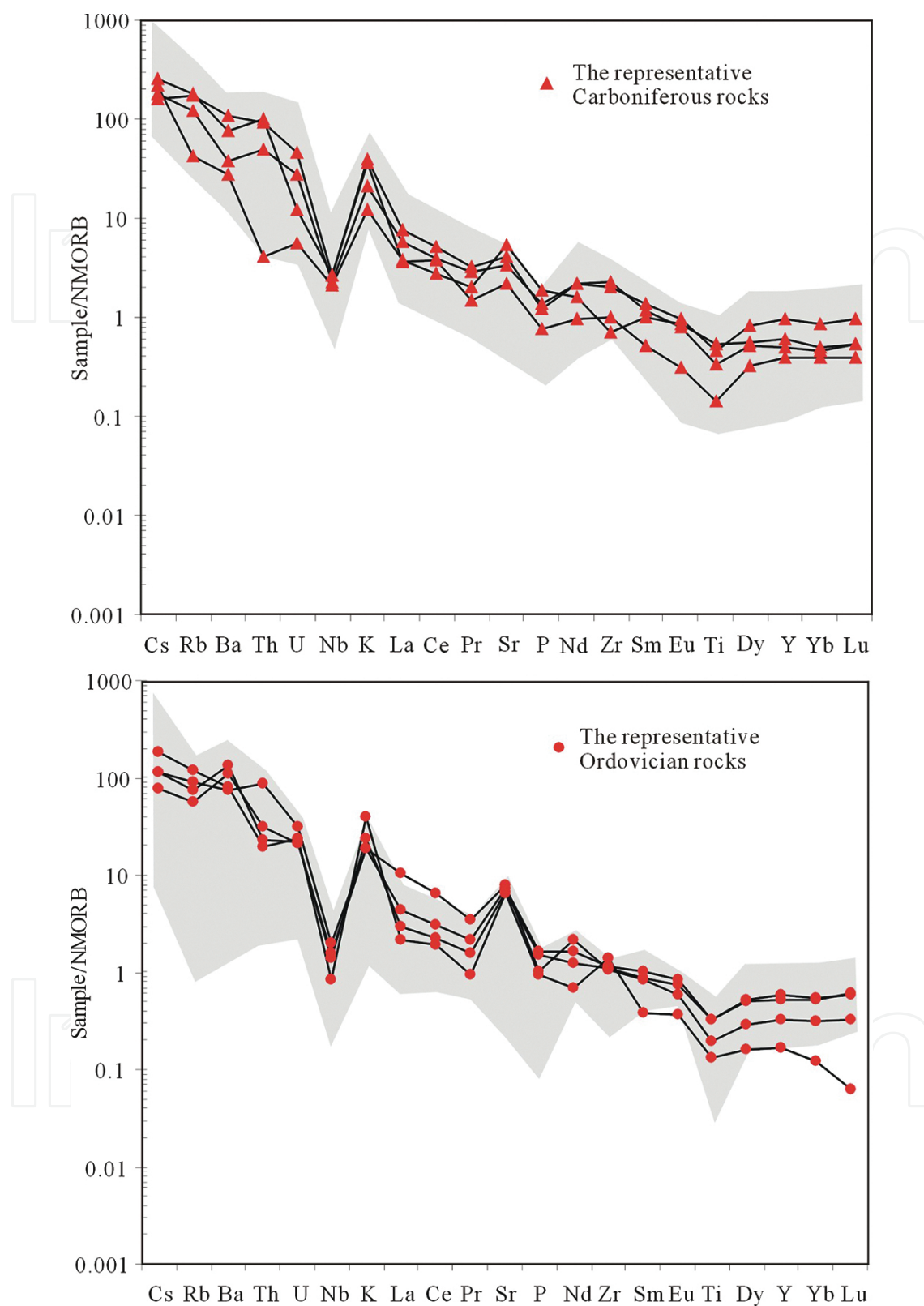


Figure 5. N-MORB-normalized trace element variation diagrams for representative samples (grey fields show data from [7, 8, 43] for Ordovician granitoids; and from [23, 30, 32, 34, 36, 37] for Carboniferous granitoids). N-MORB values are from [47].

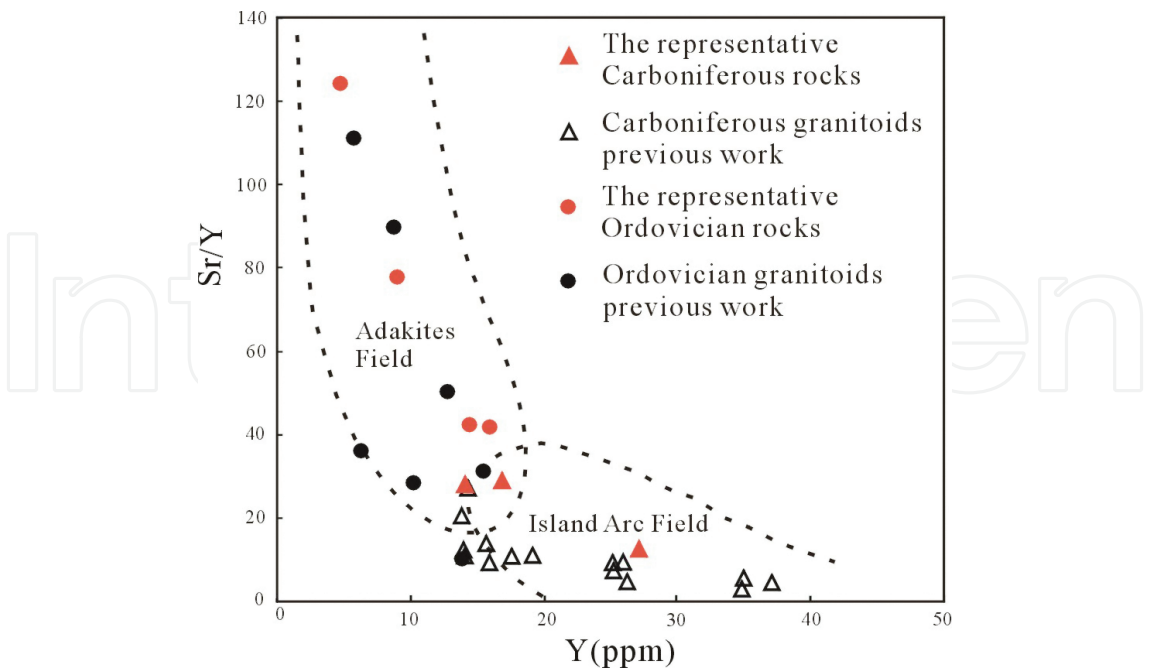


Figure 6. Y vs. Sr/Y plot showing adakitic rocks (after [18]) (data from [7, 8, 43] for Ordovician rocks; and from [23, 30, 32, 34, 36, 37] for Carboniferous rocks).

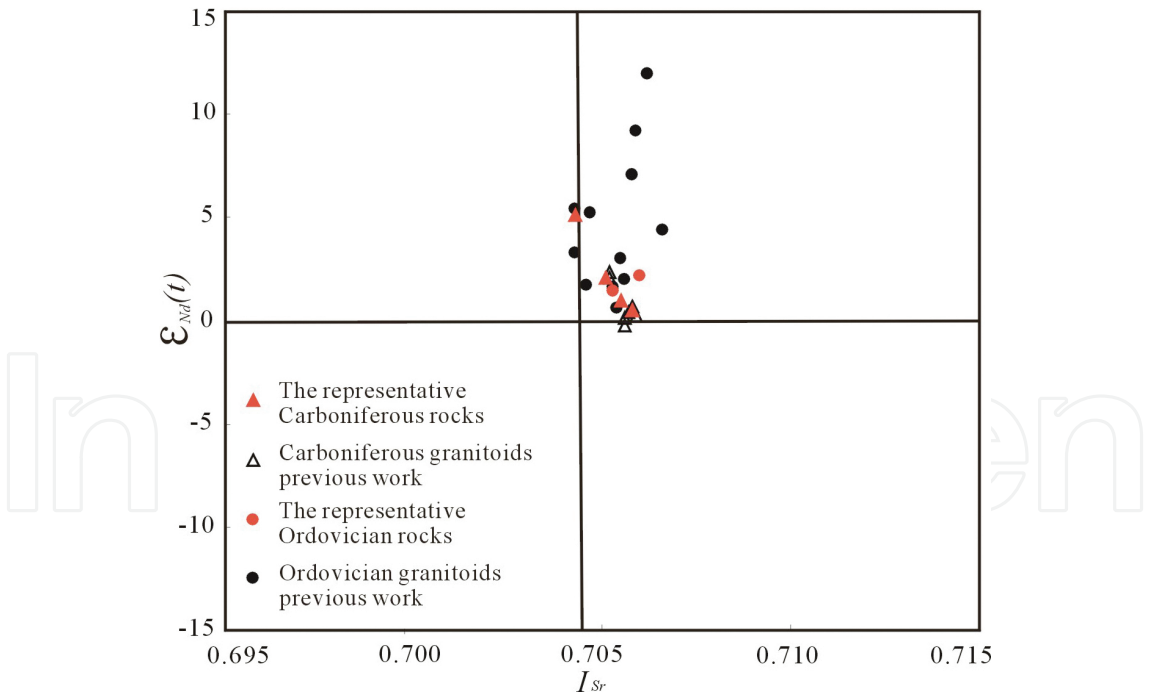


Figure 7. I_{Sr} vs. $\epsilon_{Nd(t)}$ for some typical Ordovician and Carboniferous rocks with high Sr/Y ratio from Central Inner Mongolia (data from [8, 23]).

	Ada kites			MS02-7MB1-3		MS3-5Oceanic				Active	MB1-6MB1-1		MB1-5
				Tona lite	Grano diorite	Tona lite	arc granites			continental margin arc granites	Tona arc lite	Grano diorite	Quartz- diorite
	Ada kites ^a	Cook island ^b	Cerro Pampa ^c	Oman ^d Little Port ^d				Jamaica ^d	Chile ^d				
SiO ₂ (%)	≥56.0	61.4	62.6	61.13	67.37	61.62	70.1	69.5	68.4	74.5	61.98	66.47	54.96
Al ₂ O ₃ (%)	≥15.0	18.4	17.3	17.05	15.68	16.56	12.0	14.60	14.44	12.52	16.22	15.63	18.80
Na ₂ O /K ₂ O	>1.00	7.75	3.82	2.66	3.11	1.77	15.75	4.37	1.18	0.65	2.16	1.18	3.53
MgO (%)	<3			2.34	1.14	2.27					2.95	1.69	3.65
Y (μg/g)	≤18.00	6		16.2	9.1	14.5	44	19	10	30	27.0	14.1	16.9
Yb (μg/g)	≤1.90	0.85	0.72	1.64	0.15	1.61	4.54		1.37	3.12	2.61	1.42	1.50
Sr (μg/g)	>400	1910	1886	649	711	604	200	274	210	93	304	373	473
Sr/Y	>20	319		40	78	42	4.6	14.4	21.0	3.1	11	26	28
Sr ano maly	Posi tive	Posi tive		Posi tive	Posi tive	Posi tive					Posi tive	Posi tive	Posi tive
Eu no maly	Posi tive or weakly nega tive			Weakly nega tive	Nega tive	Posi tive					Nega tive	Nega tive	Weakly posi tive
Age (Ma)	<25 Ma	<24 Ma	ca. 12 Ma	479 ± 8	472 ± 3	464 ± 8					329 ± 3	ca. 320	320 ± 3

^a [16, 49].

^b Cook island adakites [50].

^c Cerro Pampa adakites [51].

^d [52].

Table 3. The comparison of geochemical characteristics between the rocks from Central Inner Mongolia, the typical adakitic and arc rocks.

The genesis of adakites is extensively debated, and there are four proposed origins, namely partial melting of young subducted lithosphere [16], melting of newly underplated lower continental crust [53], differentiation of a parental basaltic magma [54, 55] and melting of foundered mafic lower continental crust [56]. High-Al, high Na₂O and calc-alkaline adakites are generally interpreted to have formed due to the melting of subducted oceanic crust and are different from high-K, high total alkali (Na₂O + K₂O) and low Al₂O₃ adakites that form through melting of thickened basaltic lower continental crust [16, 51, 53, 57–60].

The Inner Mongolian Ordovician granitoids of this study have depleted HREE, Nb, positive Sr anomalies, low Y and Yb contents and positive to weakly negative Eu anomalies. These characteristics are consistent with the loss of plagioclase and the presence of garnet as residual phases, probably related to partial melting of the source material under eclogite-facies conditions [61, 62]. The petrology and geochemistry of the Ordovician adakitic granitoids indicate a contribution from melting of subducted oceanic crust in their formation rather than melting of thickened basaltic lower continental crust.

The Carboniferous samples in this area have intermediate to high-silica (54.96–66.47 wt.%), high Al₂O₃ (15.63–18.80 %), higher Na₂O than K₂O (Na₂O > K₂O, Na₂O/K₂O = 1.18–3.53), low HREE (**Table 2; Figure 4**), and with low I_{Sr} (0.7043–0.7060), positive $\epsilon_{Nd}(t)$ (+1.0 to +5.1) and $\epsilon_{Hf}(t)$ (+8.1 to +12.3) isotope ratios (**Table 1; Figures 7 and 8**). However, most of them have lower Sr and Sr/Y ratio than those of Ordovician adakitic granitoids in this area (**Table 2; Figure 6**), which are typical subduction-related arc granitoids [52, 63, 64] although still with adakite-like compositions [16, 48].

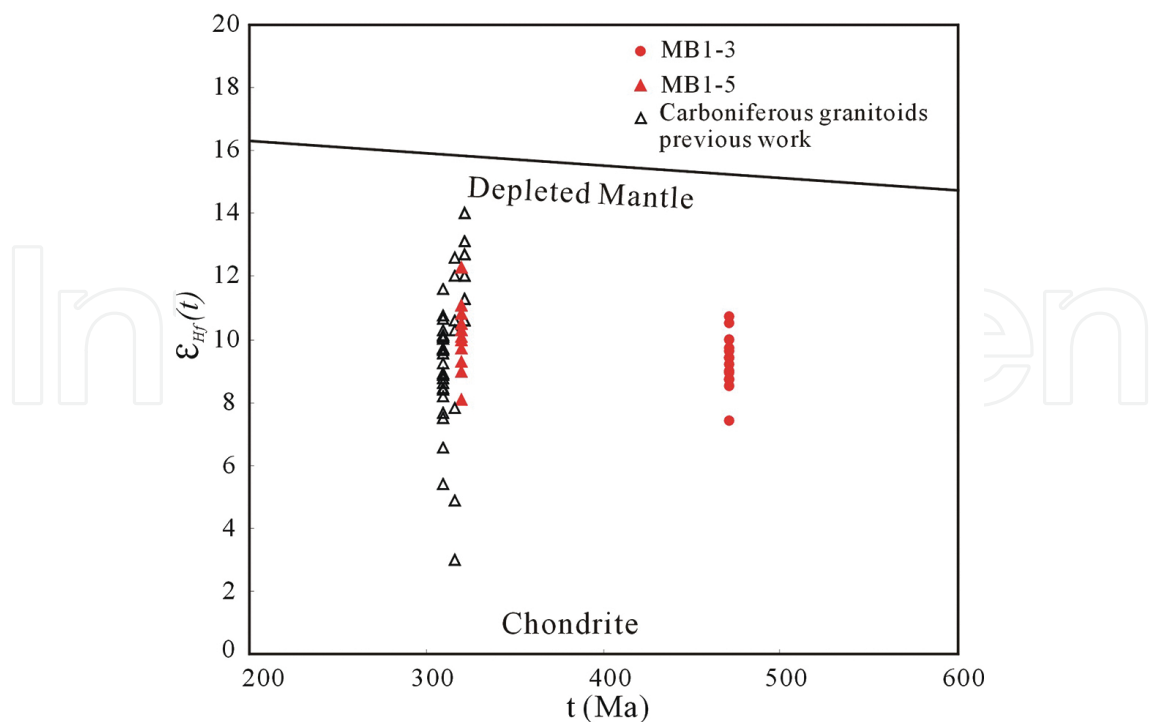


Figure 8. U-Pb age vs. $\epsilon_{Hf}(t)$ for zircons from (data from [2], and [32] for Carboniferous granitoids).

4. Geodynamic significance of the Ordovician and Carboniferous volcanic rocks and granitoids

A subduction-accretion complex usually forms along a convergent plate boundary where an oceanic plate subducts beneath another oceanic or continental plate [65]. Early Paleozoic arc plutonism as part of trench-island arc-basin systems ([6, 8, 21, 22]; **Table 2**) occurred in the southern orogen, along the northern margin of the North China Craton, and late Silurian molasse deposits unconformably overlie these rocks [6, 66]. Coeval adakitic plutonism is emplaced in the northern orogen, along the southern margin of the Mongolian terrane [20]. Silurian high-pressure metamorphic rocks [67] and Silurian syncollisional magmatism in the northern orogen along the Solonker suture [68] were also reported. All these features indicate an early-middle Paleozoic subduction and accretion system along the northern margin of the North China Craton and the southern margin of the Mongolian terrane. After demise of the ocean in the southern orogen, caused by subduction of a ridge crest and by ridge collision with supra-subduction zone ophiolite in the Silurian [8], the southern orogen became tectonically consolidated and turned into a post-orogenic setting [69].

There has been some debate about whether the Carboniferous calc-alkaline granitoids formed in a subduction zone [23, 30] or in a late- to post-orogenic setting [31]. Carboniferous calc-alkaline plutonic rocks (ca. 328–308 Ma) in the northern orogen were suggested by [2, 23, 30] as subduction genesis, which can be related to the northward subduction of Asian ocean slab. Bao et al. [31], however, thought these Carboniferous granitoids formed in a Late Paleozoic rift area because of Permian bimodal volcanic rocks. These Carboniferous granitoids include variably foliated gabbro, diorite, quartz diorite, granodiorite, tonalite and granite [23, 30], which belong to low-K tholeiitic and calc-alkaline series, and are enriched in large ion lithophile elements (LILE) and depleted in high field strength elements (HFSE) [2, 23, 30], low I_{Sr} , positive $\varepsilon_{Nd}(t)$ and $\varepsilon_{Hf}(t)$ isotope ratios ([23]) showing subduction-related arc granitoids characteristics [52, 63, 64].

Additionally, a subduction-accretion complex was identified from previously defined late Carboniferous and early Permian strata in the Daqing pasture, southern XiWuqi, Inner Mongolia [4]. In addition to this subduction-accretion complex, most magmatic rocks are considered to have formed in a subduction setting [23, 30], and the spatial configuration of both geological units indicates that the subduction polarity was from south to north [4] along the northern orogen.

Carboniferous granitoids on the northern margin of North China craton also have the composition of tholeiitic and calc-alkaline island-arc rocks and adakitic compositions [45], however, low negative whole-rock $\varepsilon_{Nd}(t)$ and zircon $\varepsilon_{Hf}(t)$ isotope ratios indicate that they were derived mainly from anatectic melting of the ancient lower crust with some involvement of mantle materials [70]. The Carboniferous plutons were interpreted as subduction-related and emplaced in an Andean-style continental-margin arc [70].

On the northern margin of the North China craton, however, Carboniferous eclogites are exposed at least 200 km south of the Solonker suture zone and have tholeiitic protoliths (MORB

and IAT), and eclogite-facies metamorphism reflects deep subduction of oceanic lithosphere [71]. The granitoids (330–298 Ma) of this area were emplaced and deformed during, and/or shortly after eclogite-facies metamorphism (ca. 331–319 Ma) [71]. This close temporal relationship indicates that magmatism closely followed the exhumation of the high-pressure metamorphic terrane [3].

5. A possible model for the discrete evolution of CAO B

The southeastern CAO B was formed by the concurrent two-way subduction of Paleo-Central Asian Ocean towards opposing continental margins in the early Paleozoic (**Figure 9A**). In the south is an arc-trench complex, which can be regarded as an analogue of the Izu-Bonin-Mariana arc [72], and in the north a product of ridge-trench interaction [8]. In the late Paleozoic, however, Andean-type orogenesis was induced by subduction of Central Asian Ocean beneath either the northern (e.g. [4]) or southern (e.g. [45]) continental blocks (**Figure 9B**). Plutonic magmatism [45] was accompanied by exhumation of a high-pressure metamorphic terrane [71] in the south; and a subduction-accretion complex [4], together with most arc-related magmatic rocks [23, 30] was formed along the northern orogen.

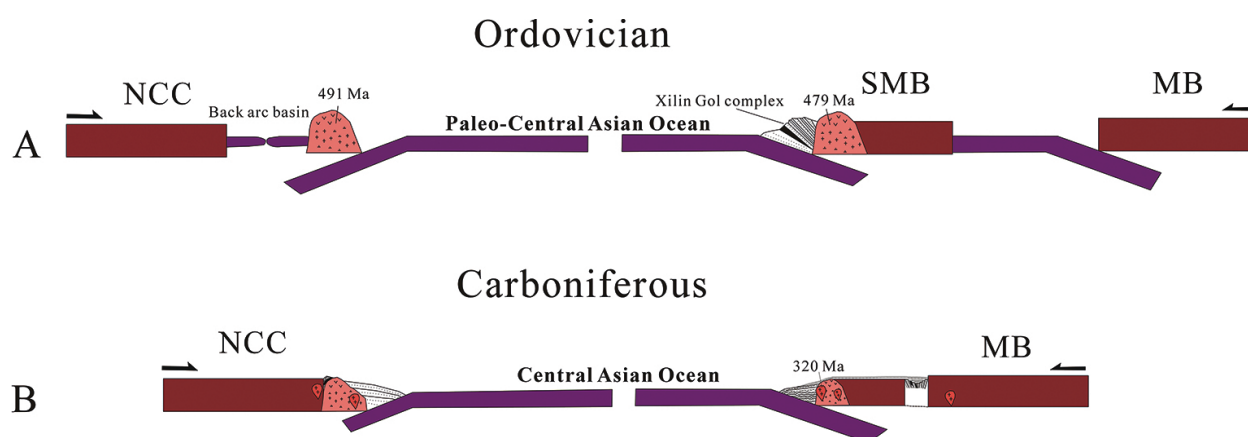


Figure 9. A possible model for Ordovician and Carboniferous evolution of Central Inner Mongolia. Abbreviation: NCC, North China Craton; SMB, South Mongolia Block; MB, Mongolia Block.

6. Modern equivalent

6.1. Cook Island and Cerro Pampa adakites

Cenozoic andesitic to dacitic rocks collected from Cerro Pampa [51] and andesites from Cook Island [50] have intermediate to high-silica, high Al_2O_3 , higher Na_2O than K_2O , low HREE,

depleted HFSE, Y and Yb, high Sr, and high Sr/Y ratios (**Table 3**), and low I_{Sr} with positive ε_{Nd} isotope ratios. The samples, therefore, represent adakites [50, 51]. Cerro Pampa adakitic magmas formed in response to melting of hot slab that was subducting beneath South America [51], and similar petrogenesis for the Austral Volcanic Zone adakites [50]. Ordovician adakitic rocks from Central Inner Mongolia show similar petrogenesis and geotectonic setting with the Cenozoic adakites from Cook Island [50], Cerro Pampa [51] and Aleutian arc [16].

6.2. Oman and Chile volcanic arc granites

Volcanic arc granites from Oman and Chile have high-silica, intermediate Al_2O_3 , low HREE [52] (**Table 3**), and with low Sr and Sr/Y ratios than the adakites (**Table 3**), which are typical subduction-related arc granitoids derived from melting in the mantle wedge. Most Carboniferous volcanic rocks and granitoids present similar petrogenesis and geotectonic setting with the Cenozoic subduction-related arc granitoids.

7. Conclusions

1. The Ordovician and Carboniferous volcanic rocks and granitoids are mostly intermediate to high-silica, high Na_2O/K_2O ratio, high Sr/Y ratios. They are characterized by LREE enrichment and exhibit typical negative Nb anomalies. All samples show positive $\varepsilon_{Hf}(t)$, $\varepsilon_{Nd}(t)$ values and low I_{Sr} .
2. The Ordovician rocks show higher Sr/Y ratio than the Carboniferous rocks, suggesting that the former represent adakitic rocks and the latter are typical subduction-related arc rocks with adakite-like compositions.
3. The Central Asian Orogenic Belt evolved through several subduction systems with different polarities in Central Inner Mongolia, namely an early-middle Paleozoic subduction and accretion system along the northern margin of the North China Craton and the southern margin of the Mongolian terrane, and late Paleozoic northward subduction along the northern orogen and exhumation of a high-pressure metamorphic terrane on the northern margin of the North China Craton.

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Author details

Yuruo Shi

Address all correspondence to: shiyuruo@bjshrimp.cn

Beijing SHRIMP Center, Institute of Geology, Chinese Academy of Geological Sciences, Beijing, China

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