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Transport of Nitrate and Ammonium During Tropical Storm and Hurricane Induced Stream Flow Events from a Southeastern USA Coastal Plain In-Stream Wetland - 1997 to 1999

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Additional information is available at the end of the chapter

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1. Introduction

Nitrogen (N) is an essential element for the functioning of aquatic ecosystems [1, 2]; however, when in excess, it can be detrimental to water quality by promoting eutrophication [3, 4] and growth of harmful microalgae and dinoflagellates [5]. Elevated N levels transported by streams and rivers into downstream coastal estuaries are of particular concern because eutrophication can affect up to 65% of the estuarine area of the coastal USA [6]. Eutrophication leads to hypoxia which stresses fish, shellfish, and invertebrates, and long-term exposure to hypoxia is fatal to most endemic fauna [7, 8]. Rapidly growing and diversifying non-point and point sources of N (e.g. agricultural crop production, urban wastewater, fertilizer use, N-enriched rainfall, and concentration of animal production systems) have been linked as causes for these troubling symptoms of eutrophication [9-12].

In eastern North Carolina, the Cape Fear River has experienced water quality issues due to high N input loads from non-point sources [13]. Mallin et al. [13] has linked water quality degradation in the Cape Fear River to N-enriched runoff from land areas with highly concentrated animal production. North Carolina has about 10 million pigs with the majority (58% of state total) of production located within the Cape Fear River basin in four counties (Duplin,

Sampson, Bladen, and Wayne) [14]. The Cape Fear and Lower Cape Fear Rivers flow through or receive input from small tributaries that drain watersheds within these four North Carolina counties. Nitrogen transported from small tributaries into the Cape Fear and Lower Cape Fear River systems has caused concern about water quality degradation [12, 15].

Non-point N movement into river and streams is a serious water quality issue in watersheds of the southeastern Coastal Plain region [16]. Nitrogen stored in soils or groundwater can move in runoff and in subsoil flow into rivers and streams especially after storm events [15] resulting in increased nutrient loading into downstream estuaries [13]. Because stream and rivers can readily transport N into estuaries, understanding processes that influence N cycling along the flow continuum offers the opportunity to shape best management strategies to reduce excess N movement. Galloway et al. [17] estimated that almost 50% of the N entering streams and rivers can be removed by N assimilative processes before it reaches coastal waters. Assimilative N processes in wetlands involves reactions between the water column and sediments [18, 19], and by microbial assimilation, denitrification and plant uptake [20, 21].

Wetlands are natural landscape features that provide an ecosystem service for N removal and are effective at reducing N loads [22, 23] and concentrations [24]. Hunt et al. [23] showed that a North Carolina Coastal Plain in-stream wetland annually removed $3 \text{ kg N ha}^{-1} \text{ d}^{-1}$ or about 37% of the total N as inflow. However, further examination of stream N export loads during direct runoff and base flow events as well as shifts between N storage pools within the wetland is needed to better understand N dynamics in this Coastal Plain ISW. It was hypothesized that greater N loads should be exported during periods of high stream outflows due to shifts in the wetland's N storage pools. This is a plausible premise because precipitation from storm events should increase stream inflows thus causing hydrologic disturbances within the wetland. In turn, these hydrologic disturbances stir up N associated with sediments, shifting equilibrium reactions between nutrients in the water column and sediment with increasing N losses with outflow. Wetland N storage/release dynamics are germane for this agriculturally intensive region of North Carolina because of the frequency of tropical storms and hurricane events [25, 26] and the subsequent increase in N movement into coastal estuaries [9, 13, 5].

The data of this study was extracted from a decades-long data base gathered from a US Department of Agricultural Water Quality Demonstration Project. The 3-year hydrological data presented in this article contained significant tropical storm and hurricane activity that is still relevant today. Current hurricane activity and tropical storms projections suggest an increase in their activity due to climate variability. Thus causing a need for past data contributions for future model projections. The objectives were to *i*) estimate annual Q_{tr} , Q_{dr} and Q_{br} and annual N (as $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$) and calculate loads exported from the ISW during these events; and *ii*) ascertain annual shifts in N storage pools between N associated with sediment, the pore water and water column.

2. Materials and methods

2.1. Watershed and In-stream Wetland Description

This study was carried out in the Herrings Marsh Run (HMR) watershed of Duplin Co., North Carolina. The HMR watershed is located in the Middle Coastal Plain region and is underlain with sandy to clay-enriched marine sediments [27]. The landscape topography is typical of the southeastern USA Coastal Plain possessing wide, nearly level to gently sloping upland areas, which have been dissected by primary, secondary and tertiary streams [27]. Soils that form in the upland areas are sandy with internal drainage ranging from somewhat poorly to excessively well-drained. Fairly wide (3-to 15-m) riparian zones form along streams resulting in floodplain areas containing very poorly drained soils [28].

For this study, daily precipitation totals were collected from both Warsaw and Clinton, North Carolina [29]. Technical information was gathered on characteristics of tropical storms and hurricanes (dates making North Carolina coastal contact, reported dates over Duplin Co., and daily precipitation) in reference [29].

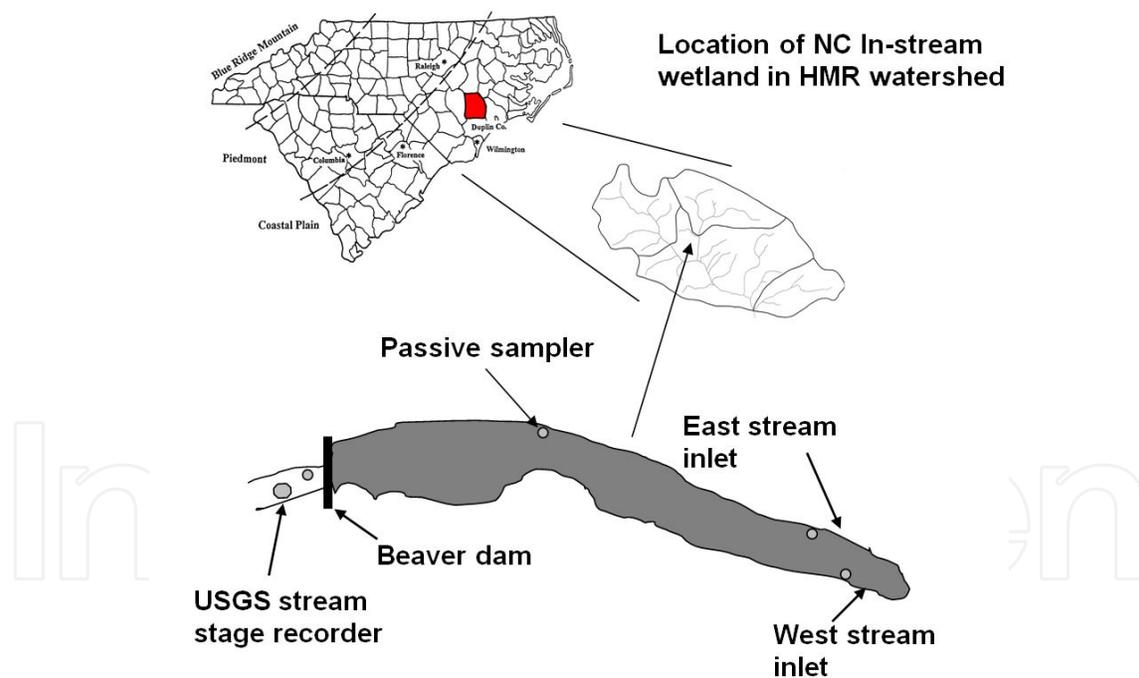


Figure 1. Location of the Herrings Marsh Run watershed in Duplin County, North Carolina. The in-stream wetland was located at the outlet of a subwatershed and equipped with an outflow stage recorder with passive samplers installed at 4 different locations (outlet, mid-point, and two inlets).

Crop and animal production practices in the HMR watershed are typical for southeastern North Carolina [28]. This includes row and truck crop production and different livestock operations. Based on a 1993 animal survey of the HMR watershed, animal heads were estimat-

ed at 29,931 for hog, 94,000 for poultry, and 176 for cattle [28]. Liquid animal manure from these operations is typically applied to row crops and pastures as a source of fertilizer.

In-stream wetlands in the HMR watershed are commonly formed in flat topographic locations and in shallow depressions along the stream continuum [27]. The studied wetland was located in the north central part of the watershed and is about 3.3 ha in size, has two diffuse stream inlets, storage volume of 29,000 m³, and the pond depth varied from < 0.3 m to almost 2 m deep (Figure 1). It receives inflowing water from two shallow second- and third-ordered black water streams on its eastern and western sides (Figure 1). This ISW is bordered by deciduous trees consisting of Bald Cypress [*Taxodium distichum* L.], Swamp Chestnut Oak [*Quercus michauxii* Nutt.], Red Maple [*Acer rubrum*, L.], Green Ash [*Fraxinus pennsylvanica* Marsh.], and both Loblolly [*Pinus taeda*, L.] and Longleaf [*Pinus palustris*, L.] pines. These trees offered a food source for colonization of the ISW by beavers (*Castor Canadensis*). Consequently, over several years, the beavers dammed up the outlet (Figure 2), which caused the two stream inflow locations to flood frequently.

2.2. Stream Sampling Locations and Stream Flow Estimates

In the early 1990s, both the east and west stream inlets flowed in well-defined stream channels. To measure water inflows, H-flumes were installed at these two inlets (Free Flow, Omaha, Nebraska⁺) equipped with pressure transducers [23]. Meanwhile, beavers improved the dam that slowed outflow and increased flooding which swamped both H- flumes. The flooding incidences lasted over several months causing inaccurate stream inflow measurements. So, there are no true inflow Q measurements available for this study.

Outflow discharges from the ISW were measured as outlined in [23] using a U.S. Geological Survey automated gauging station. Estimates of mean Q_t (as m³ d⁻¹) were obtained electronically from Q_t results that were recorded at 15-min intervals using an automated water level recorder. There were periods when daily Q_t was not recorded due to equipment damage by storms and flooding, so a linear interpolation method was used to compute values for these missing periods. The daily Q_t results were further separated into Q_{dr} and Q_b using the digital filter method [30, 31]. This digital filtering method served to partition the total flow using equations 1 to 2:

$$Q_{dr,k} = \alpha Q_{dr,k-1} + 0.5(1 + \alpha)(Q_{tk} - Q_{tk-1}) \quad (1)$$

$$Q_{b,k} = Q_{tk} - Q_{dr,k} = 0.5(Q_{tk} + Q_{tk-1}) - 0.5\alpha(Q_{tk} - Q_{tk-1}) - \alpha Q_{dr,k-1} \quad (2)$$

Where $Q_{dr,k}$ = direct runoff at time step k ; $Q_{b,k}$ = base flow at time step k ; Q_{tk} = total stream flow at time step k ; and α = filter parameter. A α of 0.925 was used in the iterations [31]. Although there are method and assumption biases in this flow separation method, Nahatan and McMahon [31] recommend the use of the digital filter method when daily Q_t results are available. Stream outflow samples were collected every 2 h using American Sigma automat-

ic water samplers (Danaher Corp., Loveland, Colorado) and later combined to make composite 3.5-d samples.

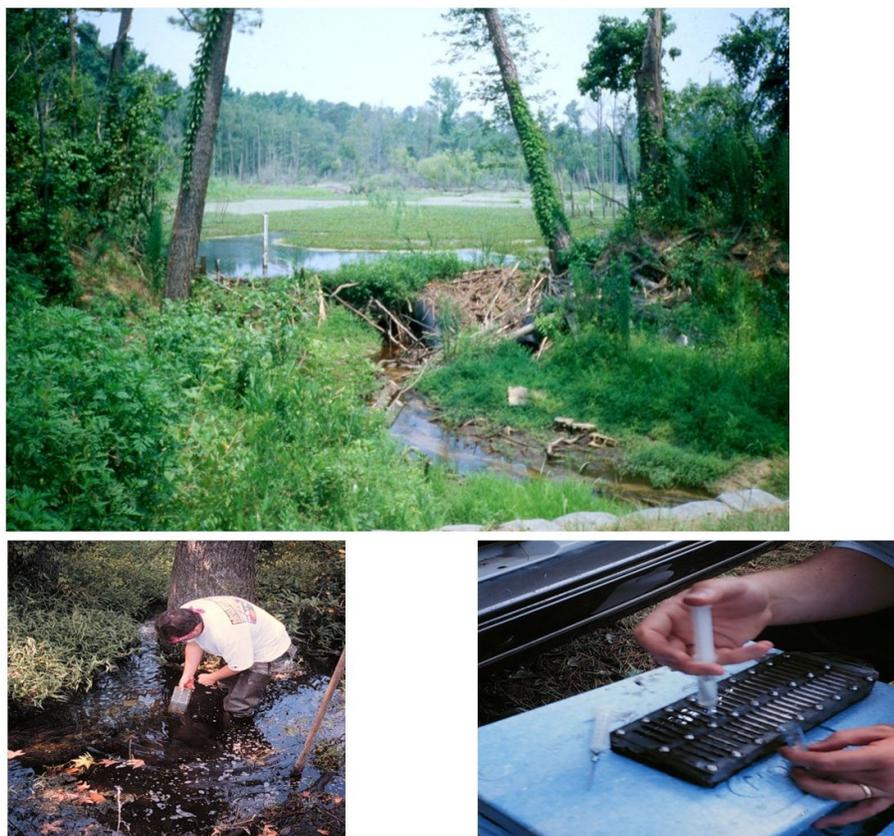


Figure 2. In-stream wetland outlet showing remnants of beaver dam (top), installation of passive samplers (bottom left), and removal of water sample using syringe from cells in passive sampler (bottom right).

2.2.1. Dissolved N Measurements

Sample preservation was accomplished by adding dilute H_2SO_4 to each automated water sampler bottle before sample collection. The acidified samples were collected weekly, filtered ($0.45\ \mu\text{m}$), and analyzed for $\text{NH}_4\text{-N}$ and $\text{NO}_3\text{-N}$ using a TRAACS 800 Auto-Analyzer (Bran + Luebbe, Buffalo Grove, Illinois) with EPA methods 353.2 and 350.1, respectively [32]. Quality control protocols for these analyses were outlined in reference [23].

2.2.2. Dissolved N Mass Loads in Wetland Outflow

From January 1997 to October 1999, daily outflow $\text{NH}_4\text{-N}$ and $\text{NO}_3\text{-N}$ mass loads (kg d^{-1}) were calculated by multiplying the daily mean flows with $\text{NH}_4\text{-N}$ and $\text{NO}_3\text{-N}$ by using their interpolated concentrations over the 3-to 4-d interval [33]. In this data set, the missing stream N concentrations were linearly interpolated to provide their daily estimated concentrations, then were used to compute corresponding daily N loads.

2.3. Passive Samplers Installation, Sediment and Water Sample Collection

Plexiglass passive samplers (peepers) were used to sample sediment pore water and the overlying water column according to procedures outlined in references [26, 33]. Each peeper consisted of a long block of plexiglass into which a series of cells spaced at 1-cm-depth increments were filled with deionized water and sealed with a membrane. Pairs of passive samplers were placed about 12-14 cm deep into the sediment at 4 locations (Figures 1 and 2) during a two-week period in Aug./Sep. during 1997, 1998 and 1999. Placement of the peepers at both ISW inlets were in shallow (0.2 to 0.4 m deep) areas, while placement at the mid-point and outlet sites was in deeper sections (0.5 to 0.8 m deep). Peepers were allowed to equilibrate with the sediment pore water and water column for 2 weeks. After removal of the peepers from the sediment, a plastic syringe was used to withdraw the sample from each cell in the peeper. The liquid sample was transferred into a plastic bottle containing dilute H_2SO_4 (preservative). The liquid samples were analyzed for $\text{NH}_4\text{-N}$ and $\text{NO}_3\text{-N}$ using the colorimetric method described earlier. Exchanges between water column and sediment N pools were determined by comparing relative differences in the water column and sediment pore water $\text{NH}_4\text{-N}$ and $\text{NO}_3\text{-N}$ concentrations. Sediment cores were also collected within 1 m of each peeper site using a bucket auger to 20 cm deep after peepers were installed. The sediments were air-dried and ground to pass a 2-mm sieve. The TKN concentrations were measured in all sediments using EPA method 351.2 [32]. Changes in the ISW N pools were determined by comparing relative differences in the annual sediment TKN and sediment pore water $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ concentrations.

2.4. Statistics

Outflow Q results, N concentrations and loads were initially computed on a daily basis and were also reported on a monthly and annual basis. This allowed for estimates of the cumulative $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ masses exported over the three-year study. To determine if the annual N loads (expressed as kg ha^{-1}) were dependent on outflow characteristics, the compiled mass of $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ by year were linearly regressed against the annual Q_t estimates. This computation was accomplished by dividing the annual N loads by the subwatershed area (425-ha). All statistical analyses were performed using SigmaStat software version 3.5 (SSPS, Chicago, Illinois).

3. Results and Discussion

3.1. Storm Activity

In the Atlantic Ocean, a distinct tropical storm (TS)/hurricane season occurs from June 1st to November 30th, with more frequent storm activity reported in August to September [34]. In this study, TS and hurricane activity was also active in August through September (Table 1). Tropical storm 1 in 1997 deposited 70 mm of precipitation over the watershed. This storm prompted an increase in Q_t from the ISW from $872 \text{ m}^3 \text{ d}^{-1}$ to $9.7 \times 10^3 \text{ m}^3 \text{ d}^{-1}$ within 24 h of the

storm's cessation. Numerous smaller storms throughout 1997 deposited less precipitation totals on this subwatershed that caused minimal impacts on Q_t . In 1998, Hurricane Bonnie deposited 183 mm of precipitation on the watershed and increased Q_t from $800 \text{ m}^3 \text{ d}^{-1}$ to $51 \times 10^3 \text{ m}^3 \text{ d}^{-1}$ within 2 days past this hurricane.

The most severe storm activity occurred in 1999. In January, three fast moving storms (FMS), which passed over the watershed in 1-2 days, collectively deposited 250 mm of precipitation. As will be presented, these three FMS abruptly influenced Q_t and $\text{NO}_3\text{-N}$ export loads. For the remainder of 1999, 1 TS and 3 successive hurricanes, which passed over the watershed, delivered impressive precipitation amounts that caused extreme hydrological disturbances to the ISW.

Month/year	Storm activity	Precipitation (mm)
August 1997	TS-1	70
August 1998	Hurricane Bonnie	183
January 1999	Fast Moving Storms (FMS)-2, 3, and 4	250
July 1999	TS-2	76
August 1999	Hurricane Dennis	166
September 1999	Hurricane Floyd	392
October 1999	Hurricane Irene	48

Table 1. Dates for tropical storms and hurricanes recorded over the study site and precipitation totals from these storm events.

3.2. Annual Flows from the In-stream Wetland

In accordance with storm activity presented in Table 1, annual Q_t flows varied considerably across three years (Table 2). The lowest annual Q_t occurred in 1997 which corresponds to the lowest annual precipitation total. The annual Q_{tr} , Q_{dr} , and Q_b for 1998 were higher than 1997 which was explained by the higher 1998 precipitation total. The highest Q_t occurred in 1999 in response to 3 FMS, 1 TS and 3 hurricane events (Dennis, Floyd, and Irene) which collectively deposited 326 and 831 mm of precipitation, respectively. In particular, the collective precipitation total from these three hurricanes accounted for 47% of the total annual 1999 precipitation total. Correspondingly, the highest annual Q_{dr} and Q_t flow from the ISW were measured in 1999, which indicates a tremendous amount of runoff entered this ISW in only a 3-month period.

Our flow monitoring results, however, showed that most water exited this ISW during base flow periods. In all three years, Q_b values from the ISW were higher than Q_{dr} with Q_b events accounting for 53 to 75% of the Q_t . The higher Q_b values implies that this ISW receives over 60% of its water supply through ground water recharge. It is not uncommon for stream Q_b to exceed Q_{dr} flows exiting watersheds in the Eastern USA Coastal Plain region [35].

Year	Precipitation (mm)	Q (m ³ × 10 ³)			Flow as a percentage of Q _t	
		Q _t	Q _{dr}	Q _b	Q _{dr}	Q _b
1997	1190	1523	554	968	36	64
1998	1410	1926	589	1337	31	69
1999	1756	8980	3226	5754	36	64

Table 2. Annual precipitation near the study area, and total (Q_t), direct runoff (Q_{dr}), and base (Q_b) flows estimated from the in-stream wetland.

3.3. Dissolved N Loads Exported from the In-Stream Wetland

3.3.1. Total Dissolved N Exports

Quantities of N exported from the ISW were expressed as a function of annual Q totals (Table 3) and as annual cumulative totals (Figure 3). The annual total quantity of both NO₃-N and NH₄-N exported from the ISW was variable. For instance, annual NO₃-N export ranged from 4910 to 24,255 kg and NH₃-N loads ranged from 778 to 1852 kg (Table 3). In 1997, more NO₃-N and NH₄-N were exported during Q_b which is consistent with the majority (64%) of ISW outflow occurring as Q_b (Table 2). The monthly amounts of NO₃-N exported increased gradually during 1997 (Figure 3). There was a pulse in NH₄-N exported during June to August 1997 in response to increased outflows from TS-1 (Table 1). In 1998, more NO₃-N was exported from the ISW than in 1997, which is reflective of the higher annual precipitation total (Table 2) and corresponding Q_t (1926 m³ × 10³). The cumulative monthly export of both NO₃-N and NH₄-N over 1998 was gradual, with no major pulses observed (Figure 3). It should be noted that there were small pulses of NO₃-N and NH₄-N exported from the ISW (144 and 37 kg, respectively) over the three days that Hurricane Bonnie delivered 183 mm of precipitation (Table 1). But, these pulses were indistinguishable in the cumulative export curves (Figure 3).

Year	Mass exported (kg)			Mass as % of total	
	Q _t	Q _{dr}	Q _b	Q _{dr}	Q _b
			NO ₃ -N		
1997	4910	1435	3475	29	71
1998	9961	2740	7176	28	72
1999	24255	7192	17063	30	70
			mean	29	71
			NH ₄ -N		
1997	778	333	445	43	57
1998	453	157	296	35	65
1999	1852	694	1158	37	63
			mean	38	62

Table 3. Annual mass estimates of dissolved NO₃-N and NH₄-N exported from the in-stream wetland during total (Q_t), direct runoff (Q_{dr}), and base (Q_b) flow events.

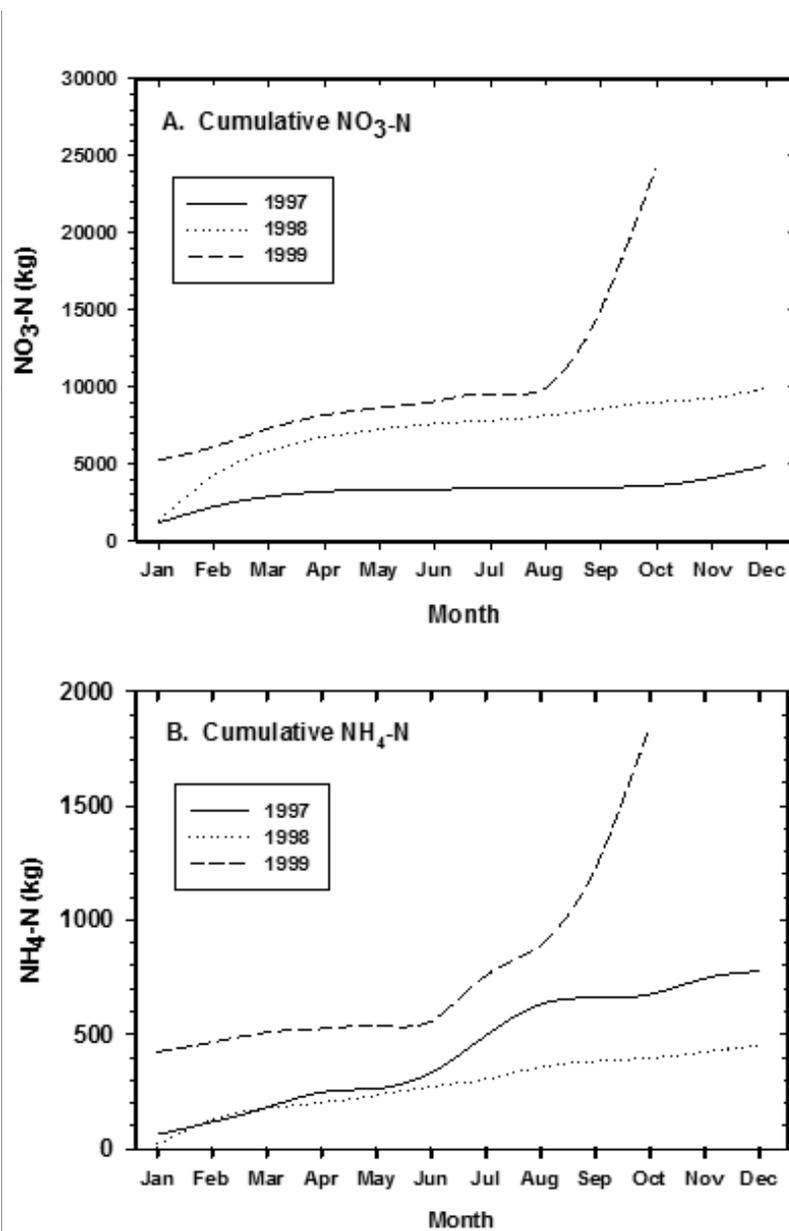


Figure 3. Cumulative annual dissolved NO₃-N (A) and NH₄-N (B) mass exported from the ISW.

3.3.2. Base Flow Dissolved N Export

In spite of the storm activity, most of the NO₃-N and NH₄-N as a percentage of the Q_t were exported from the ISW during Q_b events (Table 3). Annual base flows accounted for 70 to 72% and 57 to 65%, respectively, of NO₃-N and NH₄-N exported. Others have reported similar results that more N is exported during Q_b than Q_{dr} events [35, 36].

The greater N masses transported during Q_b events implies that most of the N entering this ISW occurred during ground water recharge. This deduction is supported in reference [37]

which reported that $\text{NO}_3\text{-N}$ discharge rates were greater for watersheds having soil or aquifer characteristics favoring $\text{NO}_3\text{-N}$ infiltration and subsurface flow. The sandy, agricultural soils in the southeastern Coastal Plain region indeed have these soil and geological features. Stone et al. [38] reported that in localized areas of the HMR watershed, farms that over-applied animal manure effluent had ground water $\text{NO}_3\text{-N}$ concentrations exceeding 20 mg L^{-1} . This supports the contention that rapid infiltration coupled with lateral flow can facilitate $\text{NO}_3\text{-N}$ leaching out of the root zone that later emerges into streams. Therefore, more $\text{NO}_3\text{-N}$ transport during stream Q_b events implies that $\text{NO}_3\text{-N}$ infiltration is a more predominant mechanism of N transport than direct runoff. The $\text{NO}_3\text{-N}$ transport mechanisms between soil and ground water, however, can be dramatically altered by storm events that favor $\text{NO}_3\text{-N}$ movement into the ISW via flooding or as direct runoff.

3.3.3. Direct Runoff Dissolved N Export

Over three years, $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ exported during Q_{dr} events from this ISW accounted for 29 to 43% of the total annual mass loads, respectively (Table 3). In 1999, however, there was a large increase in both N forms exported during Q_{dr} that prompted a closer inspection of results. The results gathered in 1999 were scrutinized to report daily precipitation totals along with a hydrograph sorted by flow events and coupled with daily $\text{NO}_3\text{-N}$ export estimates (Figure 4). Nitrate was selected for presentation instead of $\text{NH}_4\text{-N}$ simply due to the sheer mass export size differences. During January 1999, three FMS events collectively delivered 250 mm of precipitation within the subwatershed. These storms caused a moderate rise in Q_t , but also a huge increase in $\text{NO}_3\text{-N}$ was exported (5248 kg). The large spike in $\text{NO}_3\text{-N}$ load exported during January was attributed to abundant water column $\text{NO}_3\text{-N}$ concentrations (monthly mean = $6.17 \pm 1.55 \text{ mg L}^{-1}$) exiting the ISW. For comparison, monthly mean $\text{NO}_3\text{-N}$ concentrations during the warmer months of June to September were much lower at 0.84 to 3.13 mg L^{-1} , respectively. The higher $\text{NO}_3\text{-N}$ concentrations during the winter month of January are consistent with reduced N assimilation from slowed plant uptake and denitrification processes in a Coastal Plain wetland [23].

Precipitation amounts during February through July were not eventful as noted by the small spikes in the Q hydrograph and $\text{NO}_3\text{-N}$ chemograph (Figures 4 B & C). Toward the end of the traditional hurricane season, however, large amounts of precipitation (Figure 4 A, 831 mm) from the three successive hurricanes fell on the watershed causing gross hydrologic disturbances and corresponding higher Q_t (Figure 4 B). These hurricanes also induced more Q_{dr} as noted by the numerous spikes from September to October (Figure 4 B). In turn, the $\text{NO}_3\text{-N}$ chemograph responded to these hydrologic disturbances by showing dramatic increases in mass export. In fact, between October 1st and 24th, which had the subwatershed draining precipitation from Hurricanes Floyd and Irene, almost 9200 kg of $\text{NO}_3\text{-N}$ was exported from the ISW. Our results from 1999 imply that N storage by this ISW can be overwhelmed through hydrologic disturbances from storms during both non-hurricane and hurricane seasons.

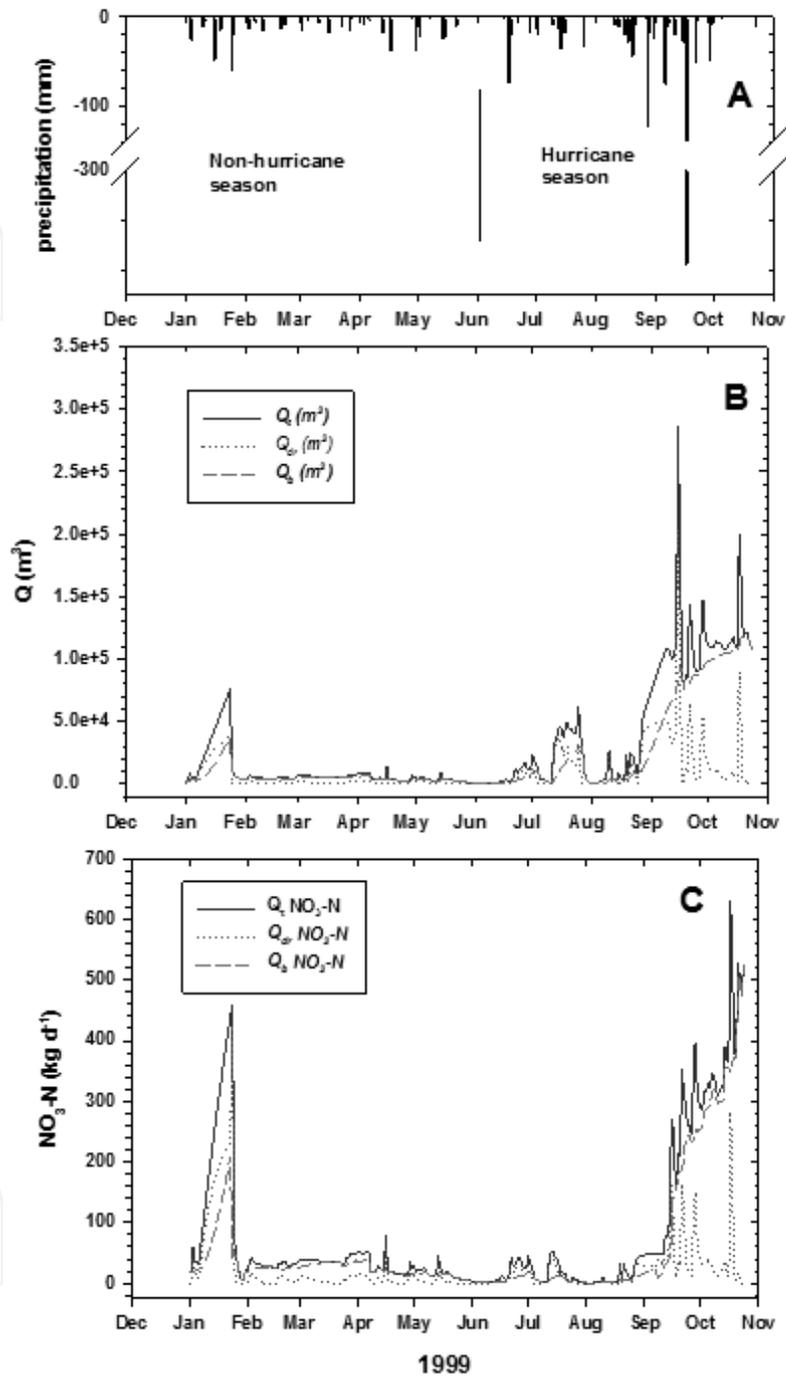


Figure 4. Daily precipitation, outflows and dissolved $\text{NO}_3\text{-N}$ exported from the ISW during 1999.

This ISW was found to release dissolved N due to storm precipitation and subsequent outflows. As more outflow occurred, correspondingly higher $\text{NO}_3\text{-N}$ and NH_4 masses were transported out. To quantify the strength of this assertion, linear regression analyses were used to determine if daily $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ mass exported from the ISW were dependent on daily Q_t (Table 4). Regression analyses revealed that the exported $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$

mass loads were significantly dependent on Q_t ; however, the strength of the predicted changes was weak to strong (r^2 between 0.21 and 0.92). There was a predictably stronger ability in 1998 to estimate daily N loads as a function of Q_t , but the strength of the relationship weakened in 1997 and 1999. Arguably, there are other processes not considered in this relationship that will influence N export loads from this ISW such as sediment associated N losses, N sequestered by biological systems, or the rate of N exchange between the water column and sediments. Nonetheless, this study demonstrates that this ISW's ability to store N forms can be compromised as outflow increases.

year	n	Slope ($\times 10^{-8}$)	Y_{int}	r^2	P
NO ₃ -N					
1997	365	278	1.84	0.40	<0.001
1998	364	510	0.21	0.92	<0.001
1999	297	224	13.92	0.62	<0.001
NH ₄ -N					
1997	365	67.5	-0.68	0.21	<0.001
1998	365	25.9	-0.13	0.71	<0.001
1999	294	16.5	1.27	0.35	<0.001

Table 4. Linear regression response between total (Q_t as L d^{-1}) outflow and pooled daily dissolved nitrate (NO₃-N) and ammonium (NH₄-N) mass loads exported (kg d^{-1}) from the in-stream wetland.

3.4. Nitrogen Storage Pools Within the In-stream Wetland

Nitrogen storage processes within a wetland involve many pathways distributed between several biotic and abiotic sources [20]. In this study, abiotic N storage pools were examined such as N associated with sediments (as TKN) and soluble N forms in the water column and sediment pore water. The relative changes in N concentrations associated with sediments and between pore water and the water column provided a snapshot of N accretion and losses. The NO₃-N and NH₄-N concentrations were measured in the water column and sediment pore water using passive samplers (peepers) at three locations within the ISW, in sediments below the beaver dam (Figure 1), and by collecting sediments near these same locations and determining their TKN concentrations (Table 5).

At both inlet locations, there were $< 11 \text{ mg L}^{-1}$ of NO₃-N in the water column and, by 1999, the concentrations declined to $< 7 \text{ mg L}^{-1}$ (Figure 5). The general decline in water column NO₃-N at both inlets was likely due to flushing of the stream sediments and riparian areas due to storm events during 1997, 1998 and 1999 [26]. The relative N concentrations between the water column and sediment pore water are important parameters to compare because they can be used to determine the gradient direction and subsequent nutrient movement response across the interface [39-41]. In all situations, sediment pore water NO₃-N concentrations at these inlets are lower than the water column, implying that the concentration

gradient would cause $\text{NO}_3\text{-N}$ movement across the sediment water column interface resulting in transfer into the sediments. The $\text{NO}_3\text{-N}$ equilibrium exchange into the sediment phase is probably accentuated through denitrification by microbial communities [19, 23]. As more $\text{NO}_3\text{-N}$ is consumed by denitrifying communities in the sediments, a concentration gradient will be created and cause its movement downward. This is a plausible explanation since the $\text{NO}_3\text{-N}$ concentrations declined rapidly as a function of sediment depth. In fact, the $\text{NO}_3\text{-N}$ concentrations declined to near zero between 1 and 3 cm deep at the East inlet (Figure 5).

At the Mid-point of the ISW, the water column $\text{NO}_3\text{-N}$ concentrations over three years showed similar decline behavior as measured at the East inlet. In that, higher $\text{NO}_3\text{-N}$ concentrations occurred in 1998 than in 1997. But both of these sites (East inlet and Mid-point) experienced a sizable $\text{NO}_3\text{-N}$ concentration decline in 1999. Similarly, $\text{NO}_3\text{-N}$ concentrations declined substantially within a few cm below the water column-sediment interface. Likewise to the $\text{NO}_3\text{-N}$ concentration gradient that occurred at both inlets, the lower sediment pore concentrations at the Mid-point over the three years would cause an equilibrium gradient shift resulting in its transfer into the sediment pore water. The downward $\text{NO}_3\text{-N}$ movement into the sediments would again cause it to be consumed by denitrifying organisms and hence its concentration declines in pore water (Figure 5).

At the outlet, water column $\text{NO}_3\text{-N}$ concentration declined to $< 0.25 \text{ mg L}^{-1}$ during 1997 and 1999, but in 1998 was as high as 3 mg L^{-1} . Sediment pore water at the outlet during the three years was almost devoid ($< 0.5 \text{ mg L}^{-1}$) of $\text{NO}_3\text{-N}$, suggesting that intensive denitrification prompted its removal from the water column.

Monitoring $\text{NO}_3\text{-N}$ concentrations in the water column and sediment pore water during the 2-week window along a flow gradient illustrated that this ISW had episodes showing outflowing $\text{NO}_3\text{-N}$ concentration reductions. Nitrate consumption by denitrifying organisms in the ISW sediments is suspected as being involved with its removal. Low outflow Q_t during the summer months could have facilitated $\text{NO}_3\text{-N}$ diffusion across the sediment interface and subsequent N removal by denitrification. This removal mechanism was probably overwhelmed by storm events that accelerated $\text{NO}_3\text{-N}$ export with outflowing water faster than could diffuse across the sediment interface and be consumed by the denitrifying microbial community.

Monitoring $\text{NH}_4\text{-N}$ concentrations using peepers along the ISW flow continuum indicated some different dynamics relative to $\text{NO}_3\text{-N}$. The $\text{NH}_4\text{-N}$ concentrations in the water column ranged from < 1 to as high as 6 mg L^{-1} (Figure 6). At all four locations, the $\text{NH}_4\text{-N}$ concentrations declined in a direction towards the sediment-water column interface. Szögi et al. [19] reported in flooded sediment, both nitrification and denitrification can occur at the same time. Similarly, the noted decline may be due to $\text{NH}_4\text{-N}$ being nitrifying by microbial communities and suggests the presence of aerobic zones above the interface allowing for organisms to conduct this oxidation [42]. The $\text{NO}_3\text{-N}$ product from this reaction could later diffuse down into the anaerobic zones in the sediments and be removed through denitrification reactions as shown by the $\text{NO}_3\text{-N}$ depletion with sediment depth (Figure 5).

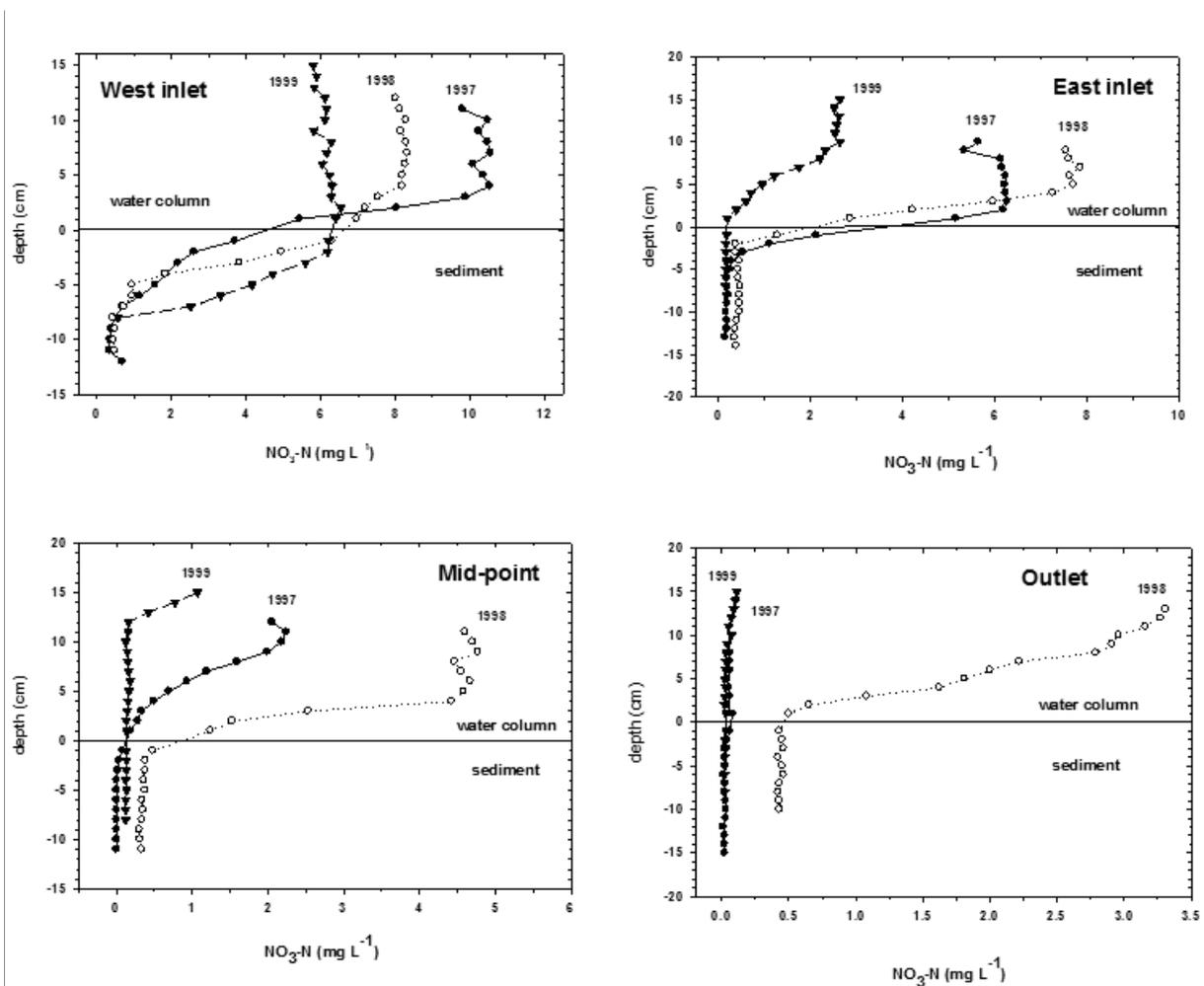


Figure 5. Nitrate ($\text{NO}_3\text{-N}$) concentrations measured in the water column and sediment at location in the in-stream wetland (1997 to 1999).

Once oxygen diffusion into the sediments becomes limited, however, the sediments will become more anaerobic thus favoring $\text{NH}_4\text{-N}$ accumulation. Ammonium does accumulate in the sediment pore water at all four locations (Figure 6). Their concentrations ranged from < 1 to almost 11 mg L^{-1} . Accumulated $\text{NH}_4\text{-N}$ in sediment pore water can move upward across the interface into the water column [19, 43]. Upward diffusion of $\text{NH}_4\text{-N}$ would occur in this system because the water columns have lower concentrations than the sediment pore water. Predictability, sediments in this ISW can act as an N source for the overlying water column because of $\text{NH}_4\text{-N}$ storage in sediments. Once diffused upward into aerobic zones, it can be nitrified to $\text{NO}_3\text{-N}$ during low flow conditions. During higher flow conditions, $\text{NO}_3\text{-N}$ would transfer with outflowing water and move down system. The balance between $\text{NH}_4\text{-N}$ upward movement, $\text{NO}_3\text{-N}$ consumption, or removal will be a function of stream flow as well as other processes (plant uptake, ammonia volatilization, etc.) known to affect N dynamics in wetlands [19, 20].

Location	1997	1998	1999
----- TKN (mg kg ⁻¹) -----			
East inlet	665	617	2536
West inlet	237	203	137
Midpoint	3371	4337	548
Outlet	854	291	322

Table 5. Sediment Total Kjeldhal Nitrogen (TKN) concentrations measured at various in-stream wetland locations.

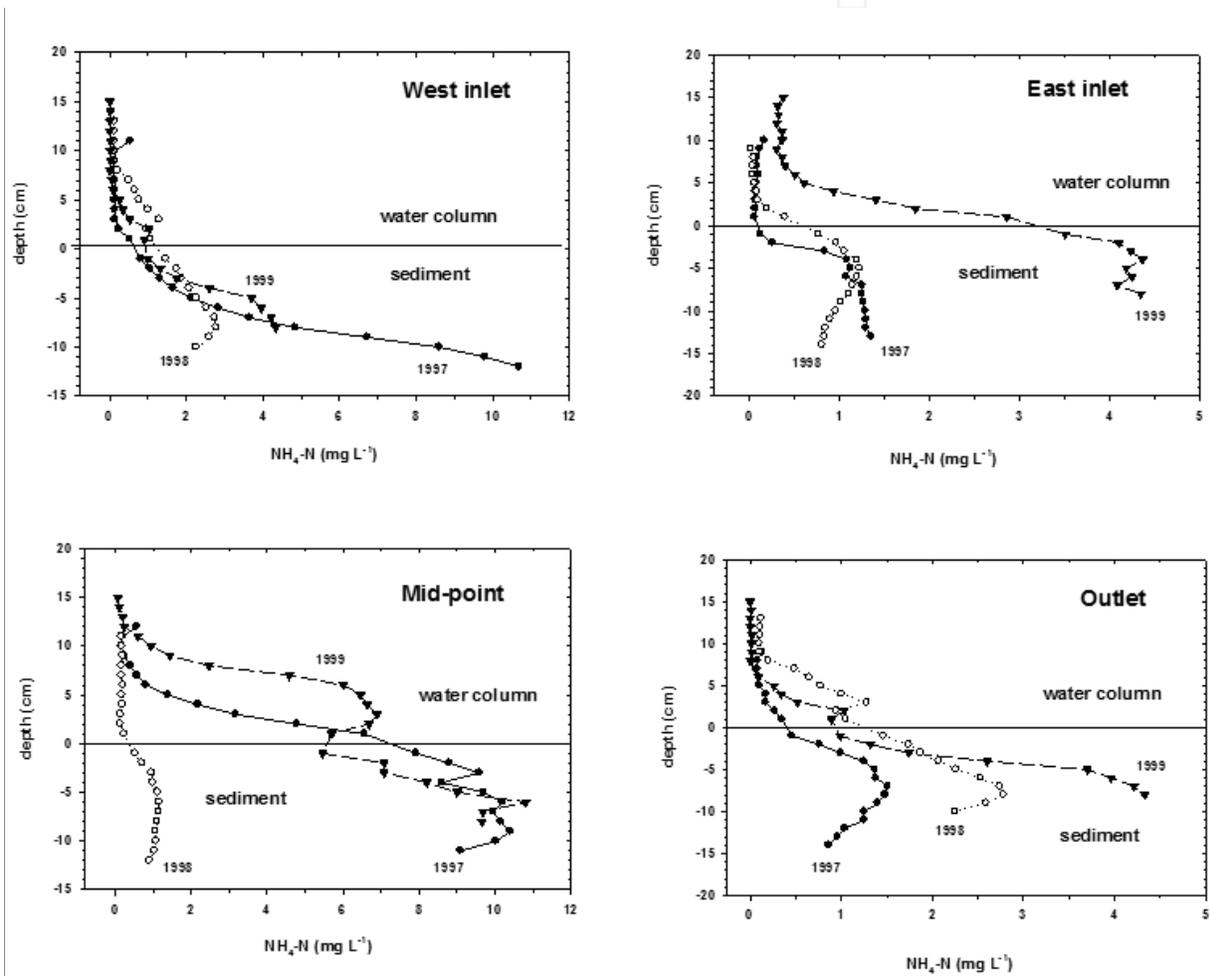


Figure 6. Ammonium ($\text{NH}_4\text{-N}$) concentrations measured in the water column and sediment at locations in the in-stream wetland (1997 to 1999).

The TKN content of the ISW sediments was measured in core samples collected when peepers were installed (Table 5). These samples provided a snap-shot of TKN contents in sediments along the flow continuum. The East inlet had higher TKN contents than the West inlet during all three sampling years. The increase in TKN content in sediments at the West inlet between 1997 and 1999 was sizable (almost 4-fold). This may be attributable to reloca-

tion of sediment enriched N from upstream riparian sources and/or N leakage from a nearby retired swine lagoon as a result of storm activity 1 to 2 months before collection [28]. On the other hand, the ISW can store N associated with sediments because some of the highest TKN contents were measured in 1997 and 1998 at its Mid-point. Between 1998 and 1999 at the Mid-point, however, there was almost an 8-fold decline in TKN content. Likewise, there was a 3-fold decline in TKN concentration at the outlet. The noted shift in sediment TKN contents implies that water flowing through the East inlet was an apparent N source for this ISW and that storm activity between 1998 and 1999 considerably reduced sediment TKN contents. As outlined [26], this ISW was inundated in 1999 with a tropical storm and Hurricane Dennis a few months before sediment collection and peeper installation. It is plausible that shifts in sediment TKN concentrations were in response to hydrologic disturbances created by these storms. As shown in Fig. 4 C, Hurricane Dennis in mid-Sep. 1999 greatly accelerated $\text{NO}_3\text{-N}$ export due to the increase in outflow (Figure 4 B).

4. Conclusions

Wetlands are important landscape features in the Southeastern USA Coastal Plain region with respect to water quality because they can act as N sinks. Their role as a natural water filter can be reversed, however, because they can act as an N source when precipitation events from tropical storms and hurricanes increase runoff and flooding causing more N transfer into the wetland. If the wetland's ability to retain this additional N is compromised, then N is released into downstream aquatic ecosystems. This study examined a North Carolina ISW's ability to retain as well as release N during different stream outflow events over a 3-yr period and also examined shifts in N storage pools using passive samplers and sediment TKN concentrations. Over 3 yr, most $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ mass loads were exported during Q_b events implying that under non-storm conditions the ISW was capable of retaining N. Results obtained in 1999, in contrast, revealed that three successive hurricane events grossly accelerated outflow Q and N releases. Not only was N export accelerated during the typical hurricane season, but a few FMS events during non-hurricane seasons also caused a large pulse of $\text{NO}_3\text{-N}$ export. During the non-hurricane seasons, this study showed that $\text{NO}_3\text{-N}$ mass loads are augmented when stream $\text{NO}_3\text{-N}$ concentrations are higher and by reduced N uptake via biological communities.

Regression analyses showed that annual $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ mass releases from this ISW were significantly correlated with Q_i flows, but the prediction of N export was weak to strong. This implies a modest ability to predict $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ export with Q_i .

As N flowed into this ISW, internal and external forces caused by chemical equilibria, flow dynamics, and biological uptake can promote shifts in the abiotic and biotic N storage pools. This study examined only abiotic N storage pools such as bound to sediments or shifted between the water column and pore water N pools. These N storage pools were subject to perturbation by external and internal dynamics forces because both sediment TKN, pore water $\text{NO}_3\text{-N}$ and $\text{NH}_4\text{-N}$ concentrations shifted in response to hydrologic disturbances from

storm events. Long-term monitoring of the ISW revealed that this ISW's ability to store/release N was dependent on outflow Q characteristics, sediment movement, and N exchanges between sediments and pore water. This ISW can store N; however, most storage occurred during low out flow (Q_b) situations because flushing of stored N in pore water and associated sediments was minimal. When outflow is accelerated, hydrologic disturbances can flush N out because of equilibria shifts between pools as well as dispersion of N-enriched sediments. It is during these large hydrologic fluctuations associated with storm events when this ISW acts as a pronounced N source for downstream aquatic ecosystems.

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